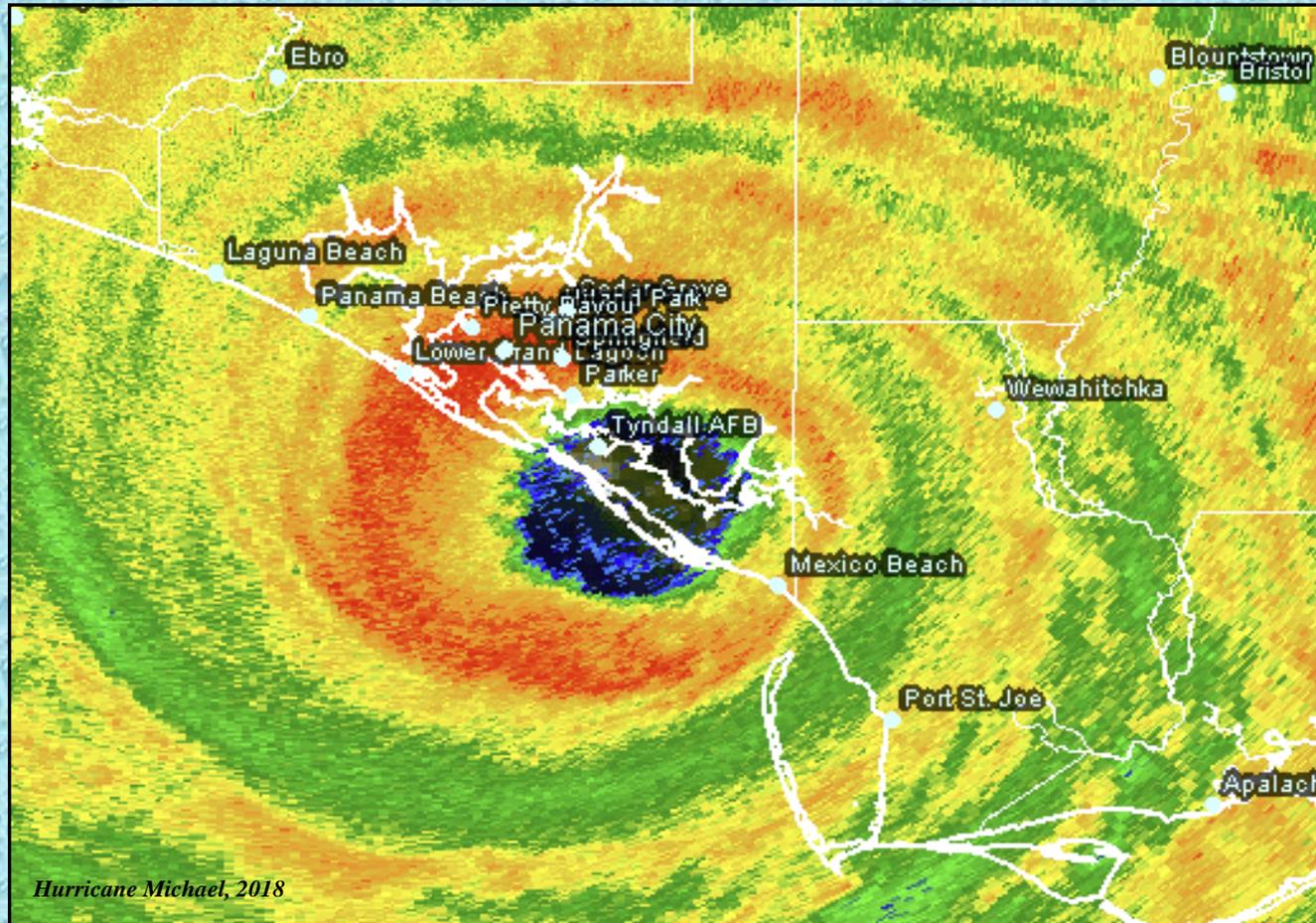


WEATHER RADAR PRINCIPLES



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6 March 2019

COURSE OBJECTIVES

Part 1. Brief Review of Basic Radar Principles

1. Wavelengths suitable for weather surveillance
2. Beam height above the surface
3. Equivalent reflectivity or dBZ
4. Z-R (Reflectivity-Rainfall) relationships
5. Doppler velocities and the '*Doppler Dilemma*'

COURSE OBJECTIVES (cont'd)

Part 2. The NOAA WSR-88D Doppler Weather Radar

1. The operational system
2. A few practical examples
3. Interpretation of the data
4. Why we need algorithms
5. A glimpse into the future

Propagation of Electromagnetic Radiation (EM)

$$V_{em} = f\lambda$$

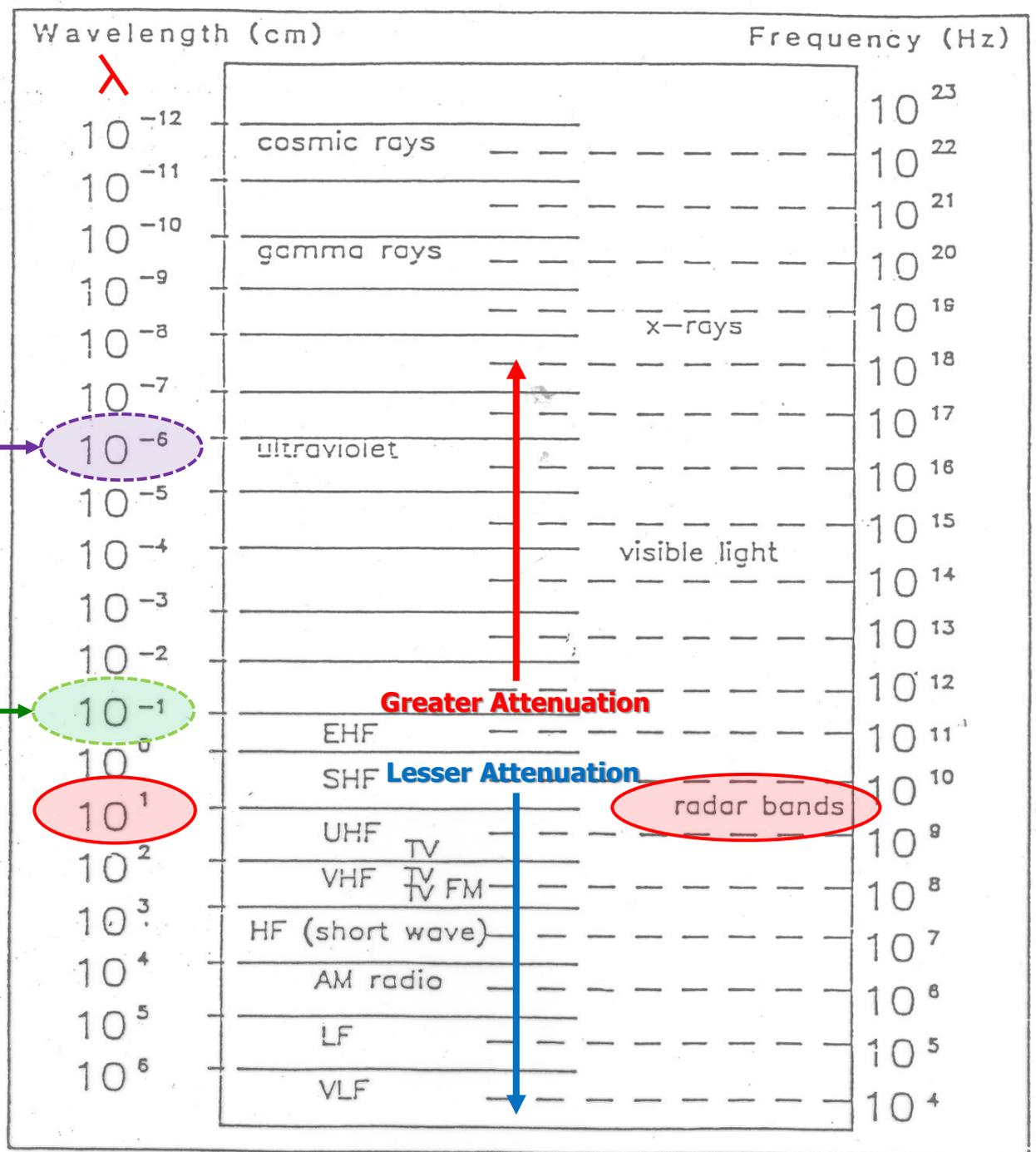
$$V_{em} \approx \text{speed of light}$$

$$= 186,000 \text{ smi/sec}$$

$$= 299,792,458 \text{ m/s}$$

Infrared satellite imagery

Passive microwave satellite imagery



Greater Attenuation

Lesser Attenuation

radar bands

Radar Operating Frequencies

| Frequency (MHz) | Wavelength (cm) | Band |
|-----------------|-----------------|-------------------------|
| 30,000..... | 1..... | K (scatterometer) |
| 10,000..... | 3..... | X |
| 6,000..... | 5..... | C |
| 3,000..... | 10..... | S |
| 1,500..... | 20..... | L (air traffic control) |

- **The longer (shorter) the wavelength, the larger (smaller) the precipitation-size particle that can be detected.**
- **The longer (shorter) the wavelength, the less (more) likely that precipitation attenuation of the radar signal will occur.**

WEATHER RADAR BANDS

10 cm

S-band

5 cm

C-band

1 cm

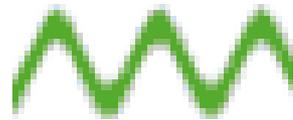
K-band

The NOAA National Weather Service WSR-88D Doppler radar is a 10-cm wavelength (S-band) weather detection radar that is excellent at sampling most precipitation particles without encountering any significant signal loss due to precipitation attenuation.

A large amount of horizontally polarized EM energy ($\sim 1,000,000$ W) is transmitted...



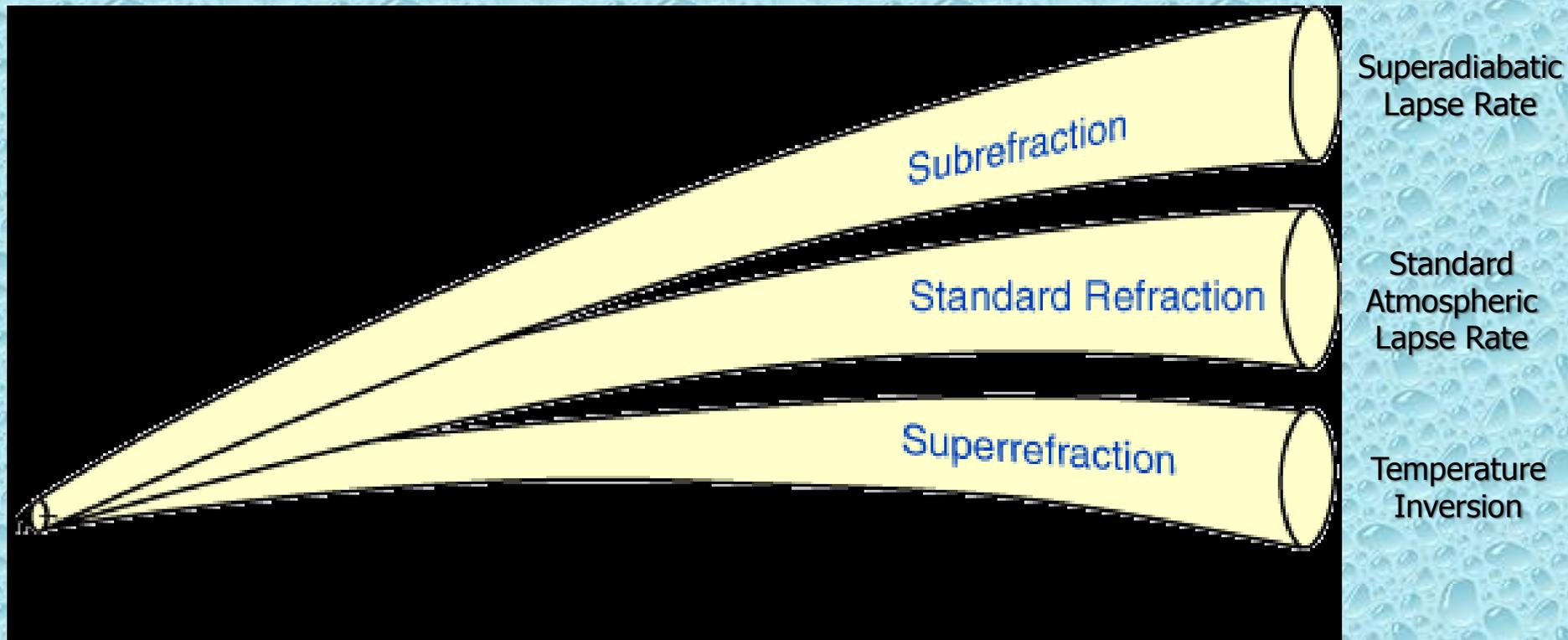
Non-isotropic (i.e., conical) radiator



Isotropic radiator

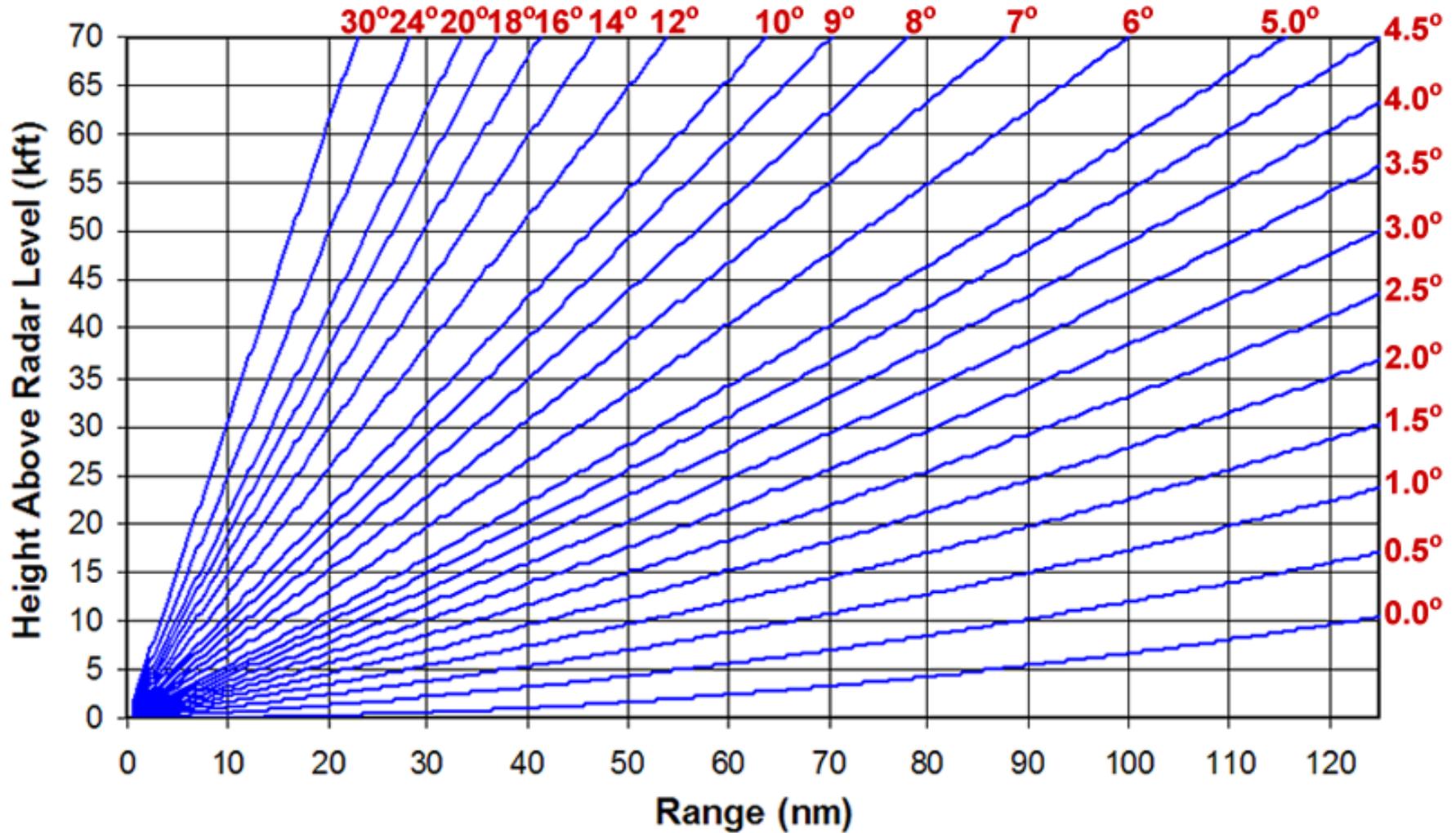
...but only a fraction of that energy (~ 0.000001 W) is 'reflected' (i.e., returned) back to the radar receiver.

Radar Beam Propagation



Differences in atmospheric density will cause the radar beam to 'bend' (i.e., refract) differently causing differences in beam centerline height from "standard".

Range vs. Height from WSR-88D Beam Height Equation



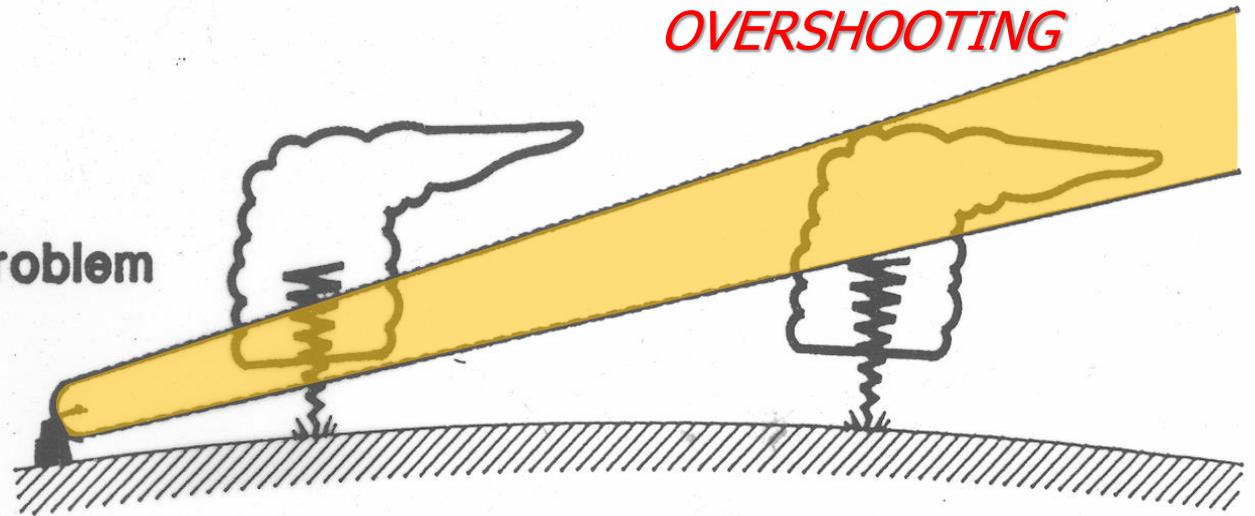
-- RHI diagrams assume standard refractivity index --

Radar Beamwidth Calculator

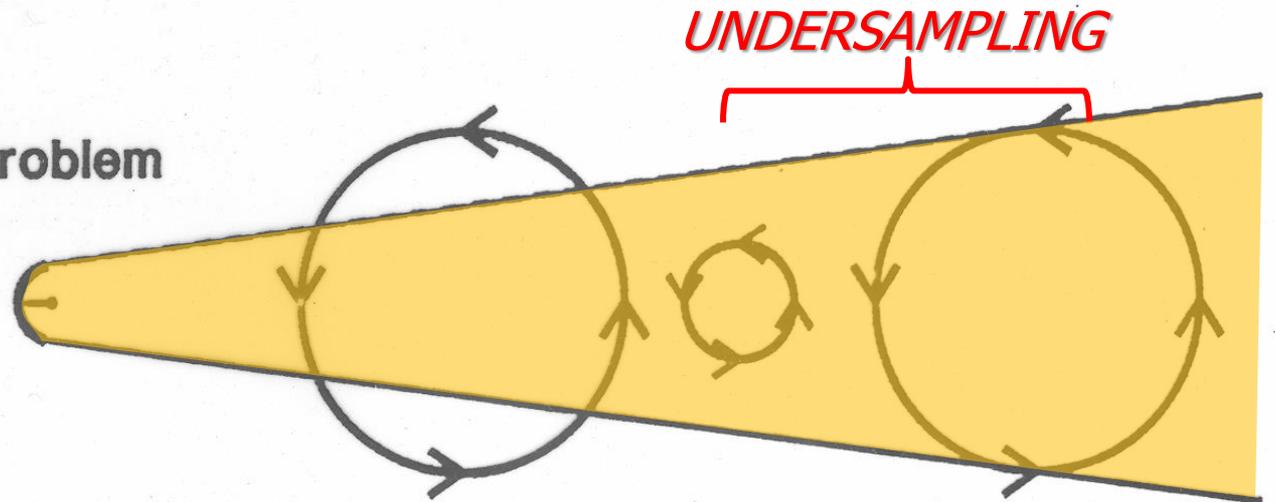
<http://www.wdtb.noaa.gov/tools/misc/beamwidth/beamwidth.html>

LIMITATIONS OF RADAR

1. Radar Horizon Problem



2. Aspect Ratio Problem



Radar Equation for Non-Isotropic Radiator

$$P_r = \frac{Z}{R^2}$$

P_r = power returned to the radar from a target (watts)

G = antenna gain

H = pulse length

K = physical constant (target character)

Z = target reflectivity

R = target range

P_t = peak transmitted power (watts)

θ = angular beamwidth

π = pi (3.141592654)

L = signal loss factors associated with attenuation and receiver detection

λ = transmitted energy wavelength

Everything inside the brackets is "known" and is, therefore, a "constant", which means that power returned to the radar by a target is directly related to the reflectivity factor, Z, and indirectly related to the range, R.

Radar Equation for Non-Isotropic Radiator

$$\overline{P_r} = \frac{P_t G^2 \theta^2 \pi^3 h |K|^2}{1024 \ln 2 R^2 \lambda^2} \sum_i D_i^6$$

$$\overline{P_r} = \frac{P_t G^2 \theta^2 \pi^3 |K|^2 Z}{1024 \ln 2 \lambda^2 R^2}$$

RETURNED POWER

Returned Power: $P_r \propto \text{Diameter}^6$

Reflectivity factor:

(for Rayleigh scattering, $D \ll \lambda$)

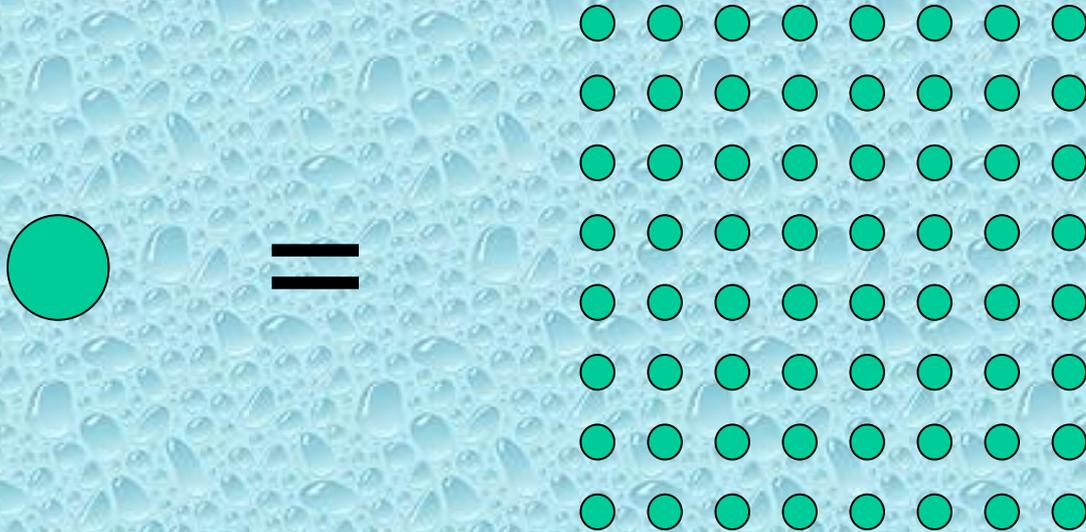
$$Z = \sum n_i \times D_i^6$$

number of
drops of
diameter D

drop
diameter(s)

- **Only a small increase in drop diameter can result in a large increase in reflectivity (Z).**
- **Large drops return the most power...but can contribute less total water mass!**

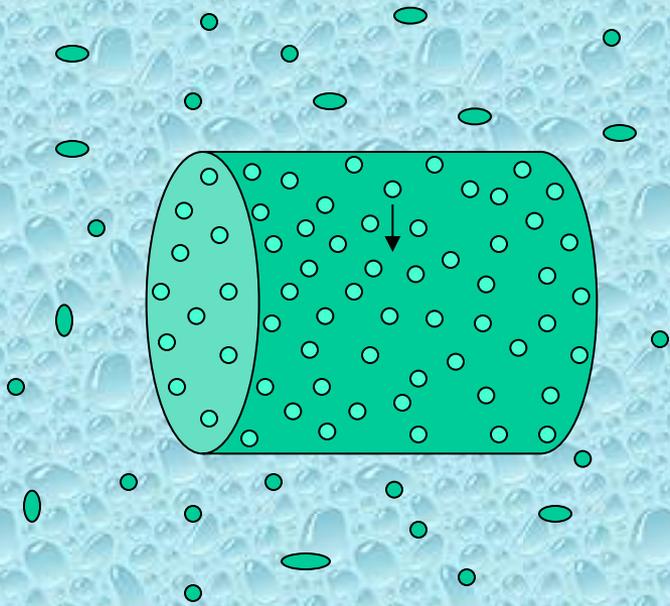
Effect of Drop Size on Reflectivity



One 1/4-inch diameter drop returns as much energy as 64 drops of 1/8-inch diameter.

However, one 1/4-inch diameter drop has a volume of only 0.065 in³, whereas sixty-four 1/8-inch diameter drops yield a volume of 0.52 in³ ...or **8 times as much total water mass!**

What would Z be for 64 drops having a diameter of only 1 mm ?

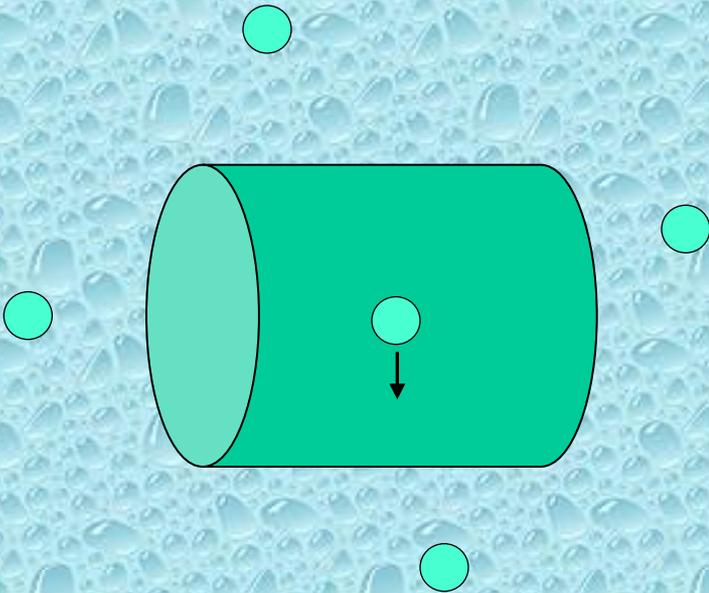


$$Z = \sum n_i \times D_i^6$$

$$Z = \sum 64 \times 1^6$$

$$Z = 64 \frac{\text{mm}^6}{\text{m}^3}$$

Now, what would Z be for only one drop having a diameter of 3 mm ?



$$Z = \sum n_i \times D_i^6$$

$$Z = \sum 1 \times 3^6$$

$$Z = 729 \frac{\text{mm}^6}{\text{m}^3}$$

REFLECTIVITY DILEMMA

The one 3-mm diameter rain drop returns more power and produces a larger reflectivity than the sixty-four 1-mm drops do... yet the one 3-mm diameter rain drop *contains less total water mass* than the sixty-four 1-mm rain drops!

Estimating Rainfall Rate Using Radar Reflectivity Data

$$Z = \sum n_i \times D_i^6$$

Since we do not know the actual drop size distribution in a radar volume sample, we use "*equivalent reflectivity*" instead of actual reflectivity.

$$Z_e = \frac{P_r \times R^2}{const}$$

where, P_r = power returned
 R = target range

Equivalent reflectivity

$$dBZ_e = 10 \times \log(Z_e)$$

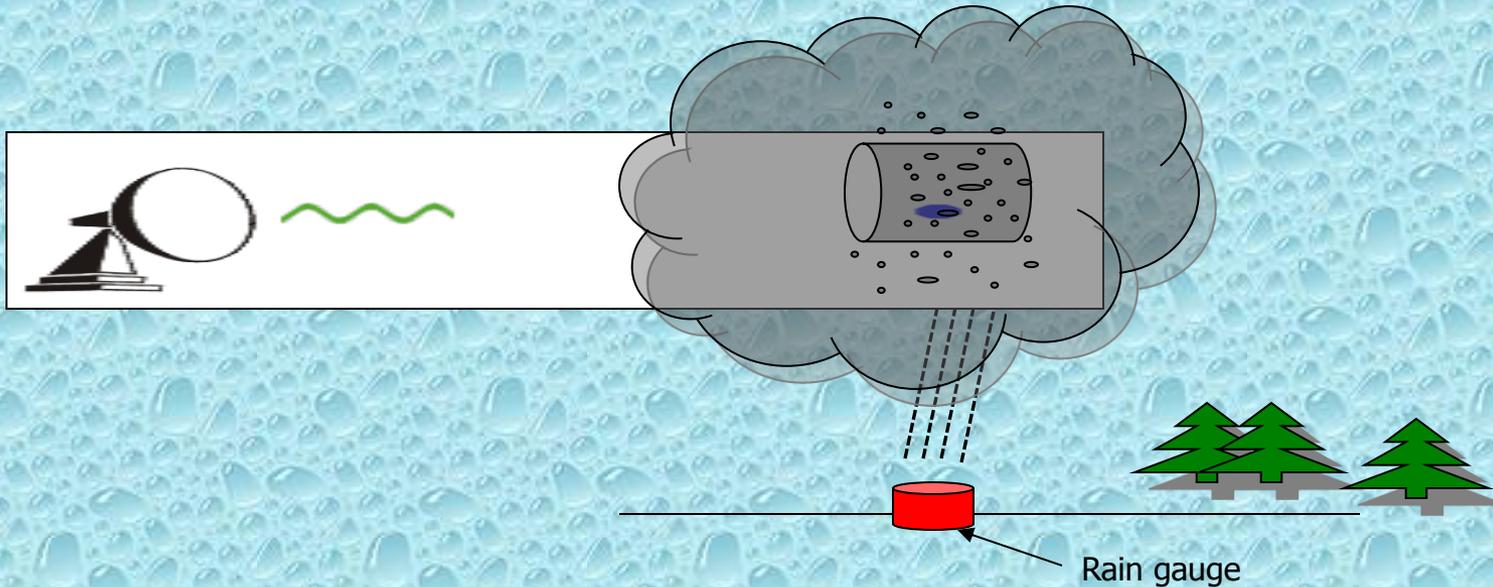
Using 10 times the logarithm of Z_e keeps the range of values of Z_e small, but still operationally useful.

$$dBZ_e = 10 \times \log(Z_e)$$

| Z_e | Log Z_e | dBZ_e |
|------------|-----------|---------|
| 10 | 1 | 10 |
| 100 | 2 | 20 |
| 1,000 | 3 | 30 |
| 10,000 | 4 | 40 |
| 100,000 | 5 | 50 |
| 1,000,000 | 6 | 60 |
| 10,000,000 | 7 | 70 |

Z-R or Reflectivity-Rainfall Relationships

↑
we now have the input we need (i.e. Z_e)



Find an empirical relationship to estimate rainfall rate:

$$Z_e = a R^b$$

$$Z_e = 300 R^{1.4}$$

Rainfall Rates (in mm hr^{-1}) for Various Z - R Relationships

| | WSR-88D | Conventional | Convective | Snowfall |
|-----------|----------------|--|---------------|-----------|
| dBZ | $300R^{1.4}$ | $200R^{1.6}$ | $486R^{1.37}$ | $2000R^2$ |
| 20 | 0.02\0.05 | 0.03\0.76 | 0.01\0.25 | 0.01\0.25 |
| 30 | 0.09\2.28 | 0.12\3.05 | 0.07\1.78 | 0.03\0.76 |
| 40 | 0.48\12.2 | 0.47\11.9 | 0.36\9.14 | 0.09\2.29 |
| 50 | 2.50\63.5 | 1.90\48.3 | 1.90\48.3 | 0.28\7.11 |
| 55 | 5.7\145 | (55 dBZ = maximum reflectivity used for rainfall conversion by WSR-88D) | | |
| 60 | 12.9\327 | 8.10\306 | 10.3\262 | 0.88\22.4 |
| 70 | 67.0\1702 | 34.1\866 | 55.4\1407 | 2.78\70.7 |

Probable
Wet Hail
Contamination



Rainfall Rates (in mm hr⁻¹) for WSR-88D Tropical Z-R Relationship

minimum radar reflectivity for determining eyewall diameter

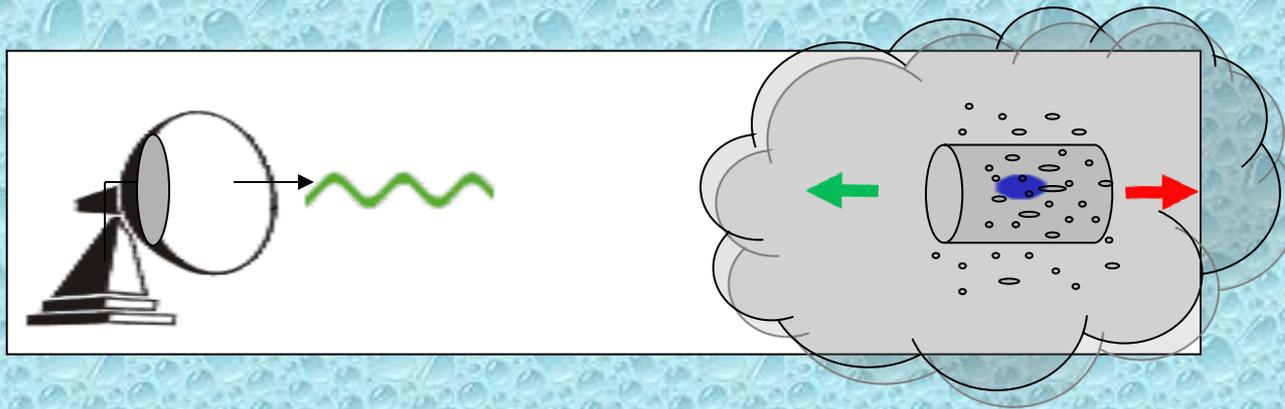
| dBZ | Z | 250R ^{1.2} |
|-----|----------|---------------------|
| 15 | 31.6 | 0.01\0.18 |
| 20 | 100.0 | 0.02\0.47 |
| 25 | 316.2 | 0.05\1.22 |
| 30 | 1000.0 | 0.12\3.17 |
| 35 | 3162.3 | 0.33\8.28 |
| 40 | 10000.0 | 0.85\21.6 |
| 45 | 31622.8 | 2.22\56.5 |
| 50 | 100000.0 | 5.80\147.4 |
| 55 | 316227.8 | 15.14\384.6 |

$$R = \sqrt[1.2]{\frac{Z}{250}}$$

**Radar Detection of
Atmospheric Motion**

or

Doppler Velocities



In addition to a measurement of power (reflectivity), we also have a measurement of particle motion.

A Doppler weather radar measures a single component of motion, but only **toward** or **away** from the radar.

The “Doppler Dilemma”

1. Speed of light c
2. Wavelength λ
3. PRF (pulse repetition frequency)

$$R_{\max} = \frac{c}{2PRF}$$

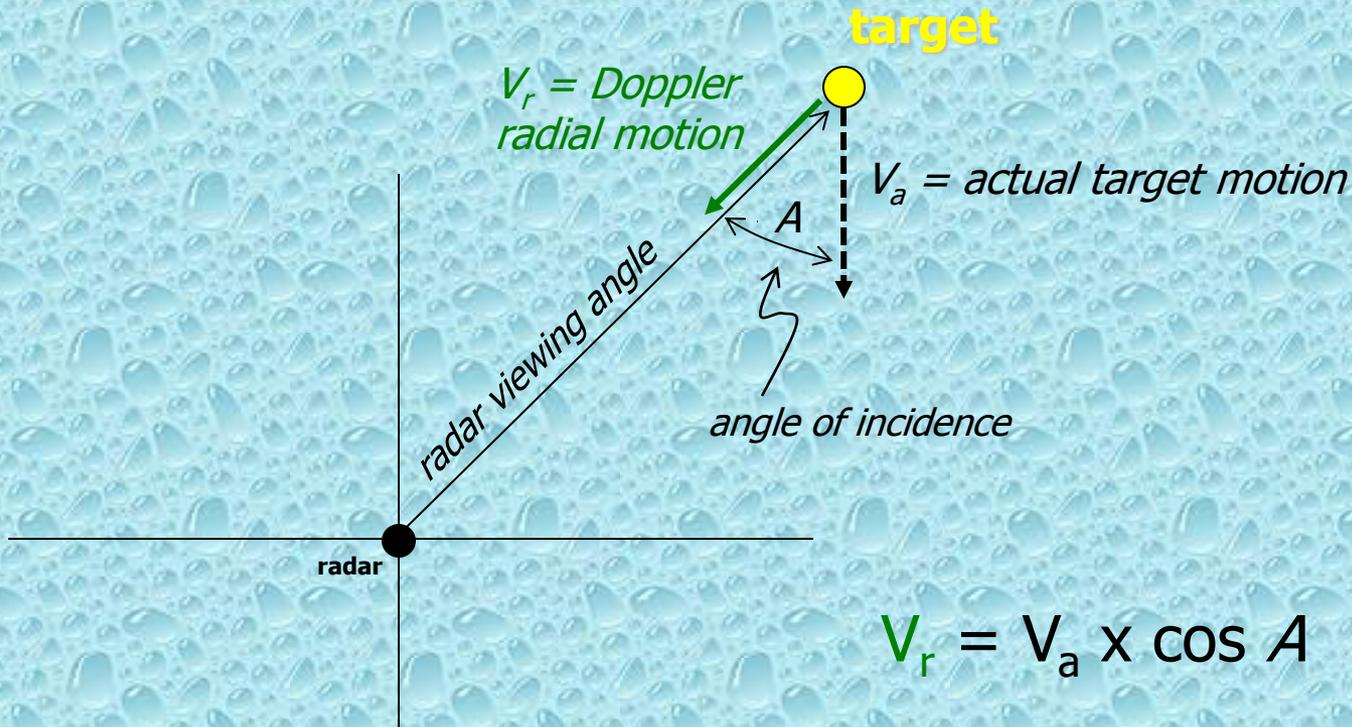
Maximum Unambiguous Range

but,

$$V_{\max} = PRF \frac{\lambda}{4}$$

Maximum Unambiguous Velocity

Example of Actual Velocity => $V_a = 20$ kt



$$\begin{aligned} V_r &= V_a \times \cos A \\ &= 20 \text{ kt} \times \cos 45^\circ \\ &= 20 \times .707 \end{aligned}$$

$$V_r = 14.14 \text{ kt}$$

Part 2

NOAA WSR-88D Doppler Weather Radar

- Weather Surveillance Radar 1988-Doppler
- first working prototype installed in Norman, OK in 1988
- This is the radar used operationally by the U.S. National Weather Service
- NEXRAD is the name of the federal procurement program which developed the WSR-88D

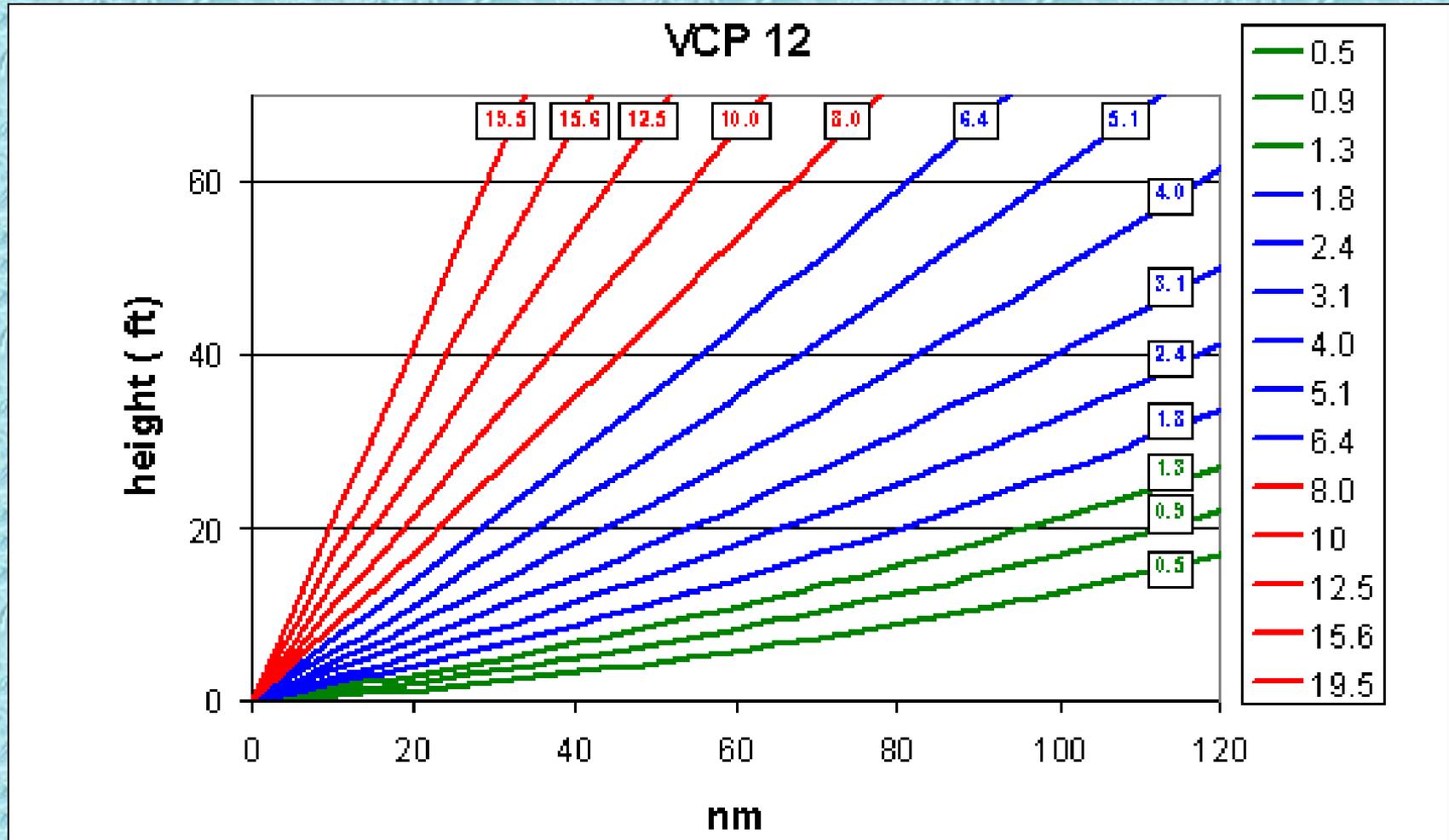




Andrew Hatzos NWS Cincinnati OH



WSR-88D radar utilizes 9 different Volume Scans to collect reflectivity and Doppler velocity data



Hurricane Lenny, November 1999



BASE
REFLECTIVITY

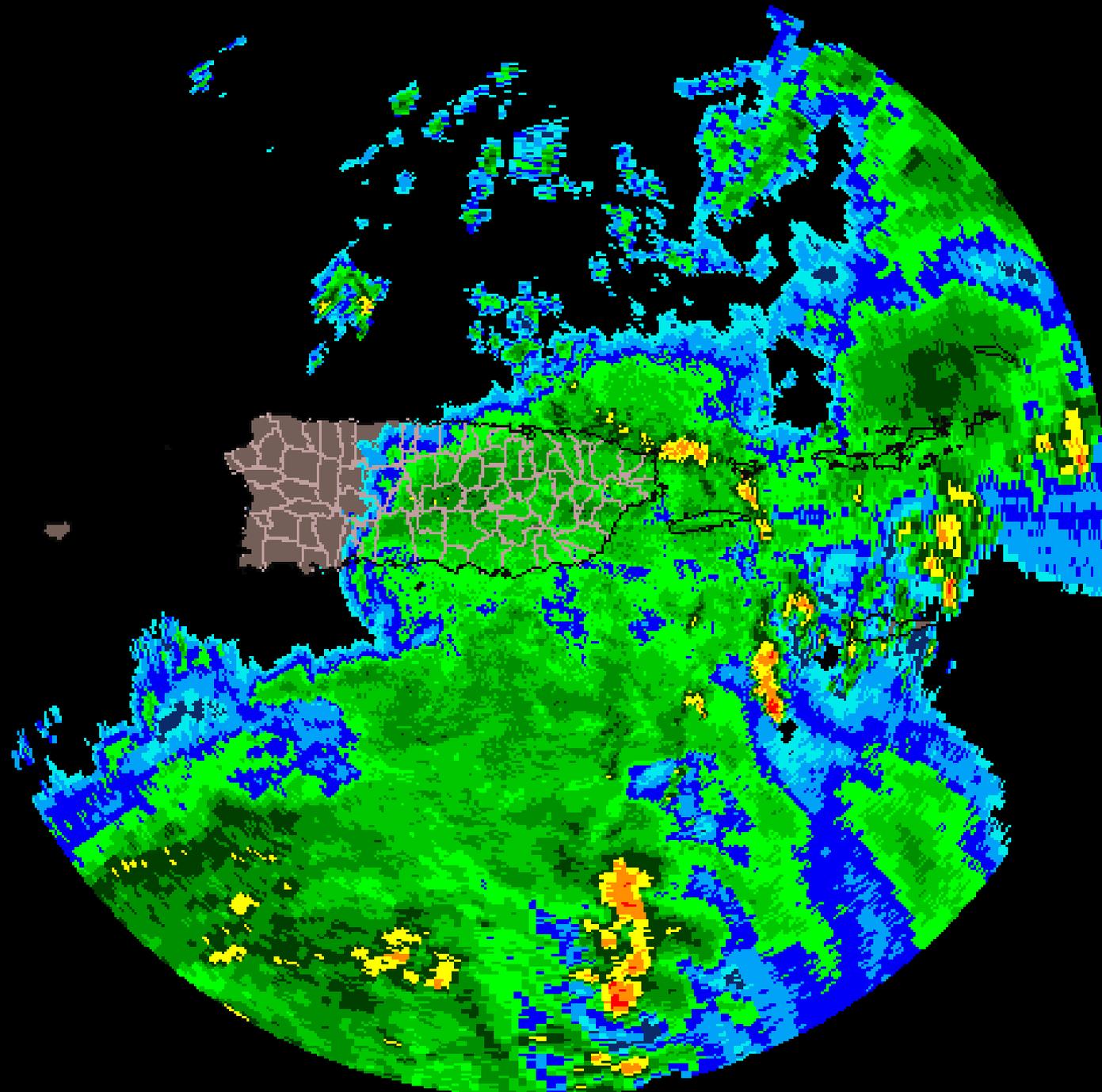
JUA

11/17/99 0112Z
RANGE: 230 KM
RES: 1 KM X 1 DEGREE
MODE: PRECIPITATION
ELEV: 0.5 DEGREES

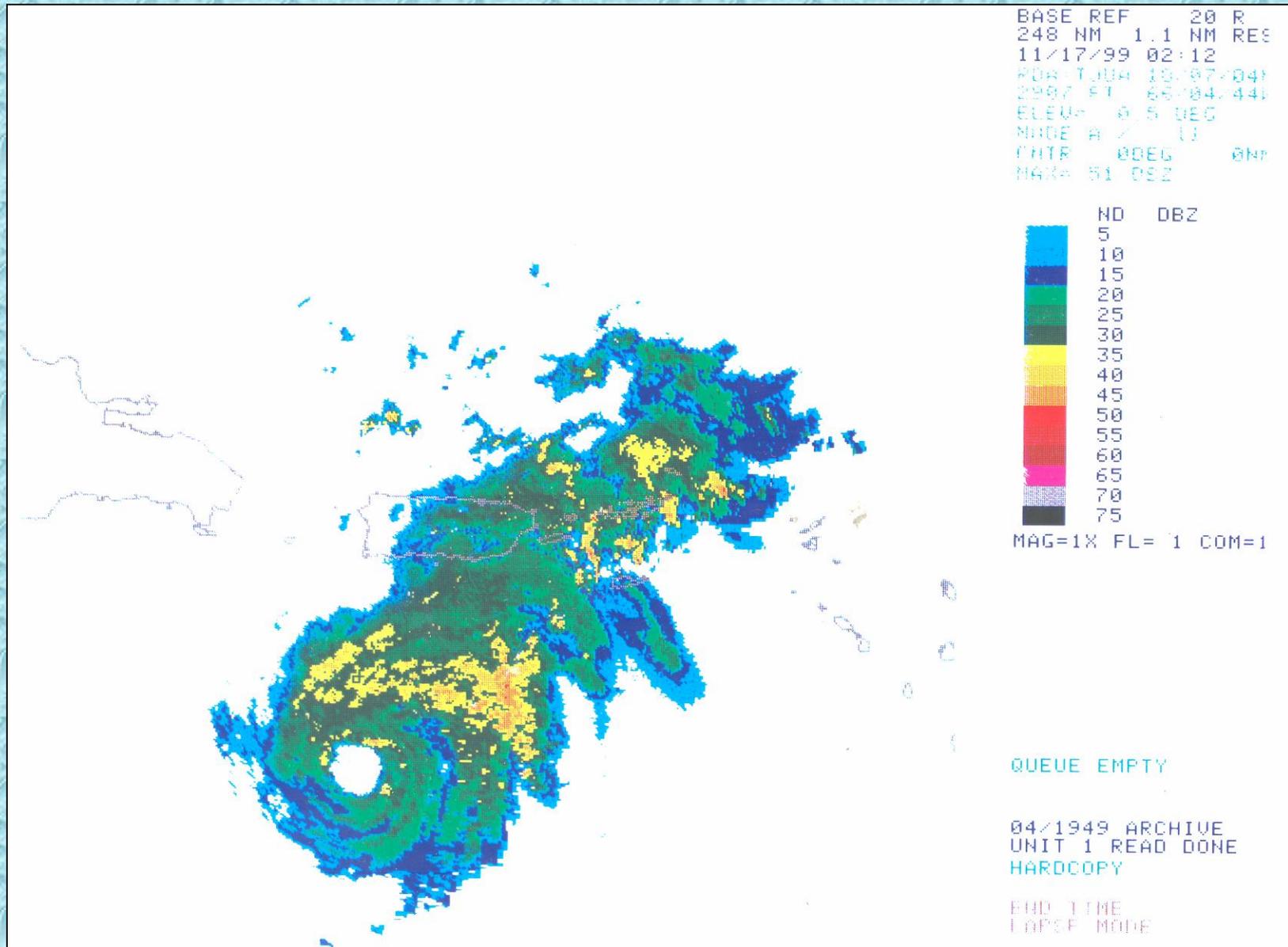
DBZ



MAX DBZ: 53



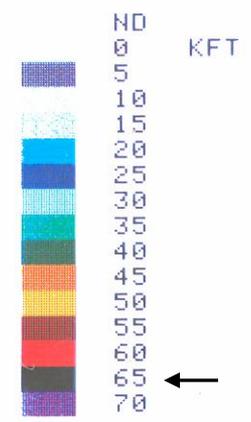
Hurricane Lenny (1999) – 100 kt intensity at 0212 UTC



"Wrong-Way" Lenny (1999)

12/04/99 20:05
ECHO TOPS 41 ET
124 NM 2.2 NM RES
11/17/99 02:12
EQA T.00A 18.87/94
2497 FT 66.84/445

MODE A 7 11
CNTR 00EG 0NF
MAX= 66 KFT

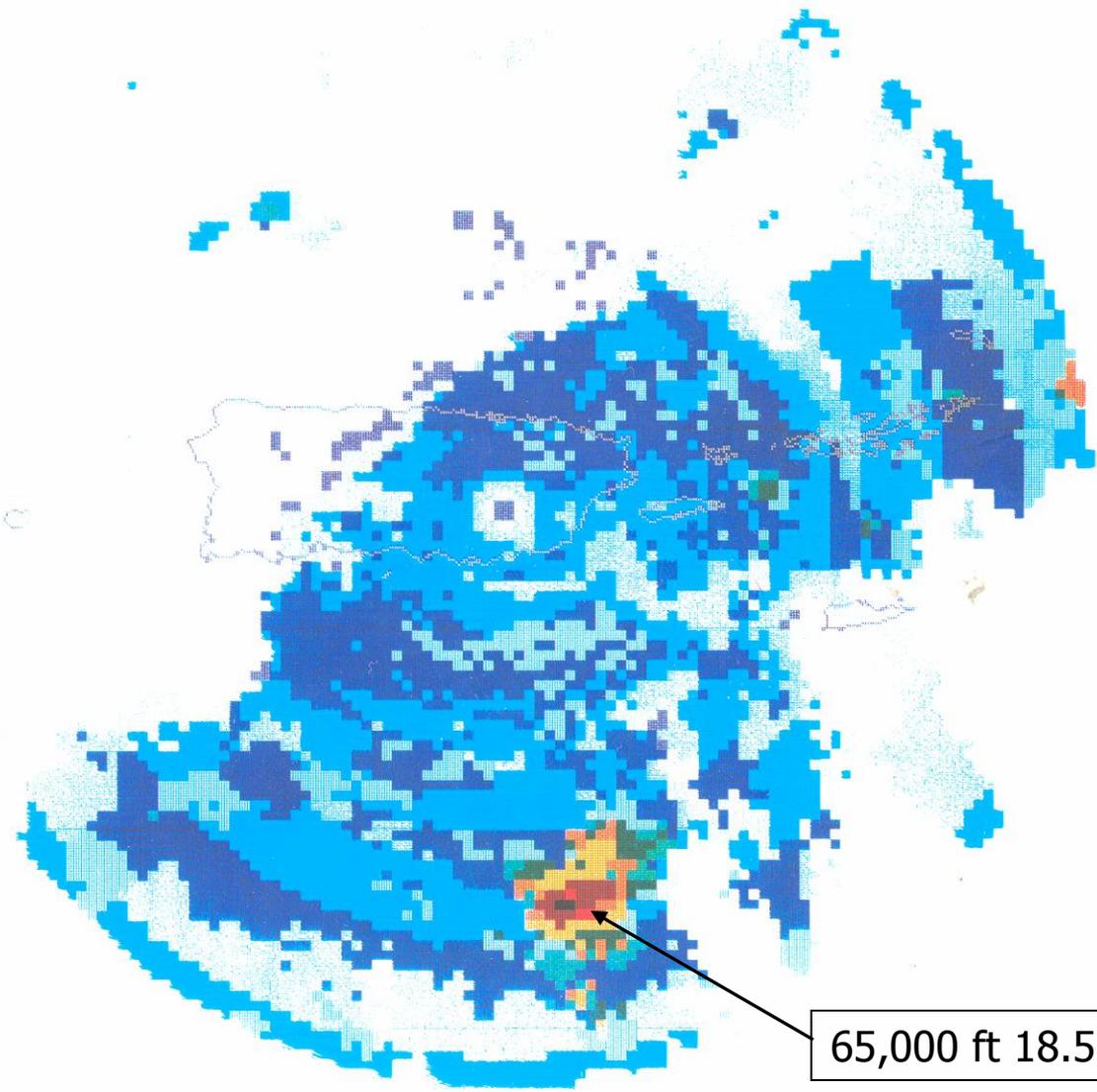


MAG=1X FL= 1 COM=1

TL 3 RATE= 1 0 SEC
A-R (PGA)

QUEUE EMPTY

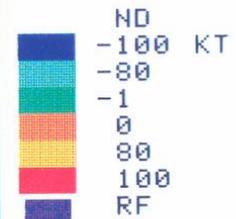
04/1949 ARCHIVE
UNIT 1 READ DONE
HARDCOPY



65,000 ft 18.5 dBZ echo top!

...2 hours later...

BASE VEL 24 U
124 NM .54 NM RES
11/17/99 04:13
RDA:TJUA 18/07/04N
2907 FT 66/04/44W
ELEV= 0.5 DEG
MODE A / 11
CNTR 165DEG 51NM
MAX=-101 KT 49 KT



MAG=4X FL= 1 COM=1

A/R (RDA)

QUEUE EMPTY

04/2120 ARCHIVE
UNIT 1 READ DONE
HARDCOPY

Hurricane Lenny with 105-kt intensity at 0413 UTC -- but why did the radar only show 80-99 kt Doppler velocities?

BASE
REFLECTIVITY

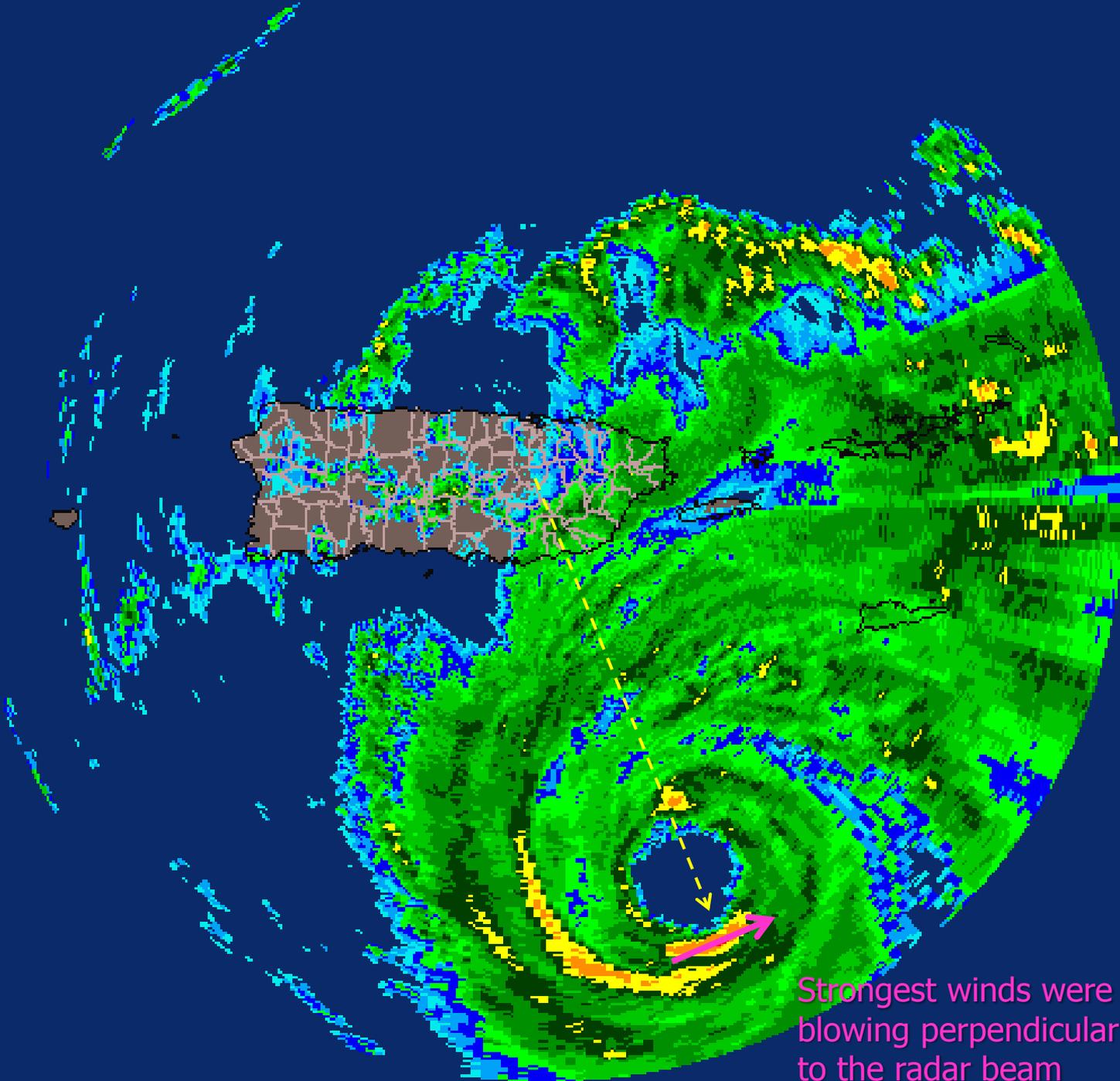
JUA

11/17/99 1141Z
RANGE: 230 KM
RES: 1 KM X 1 DEGREE
MODE: PRECIPITATION
ELEV: 0.5 DEGREES

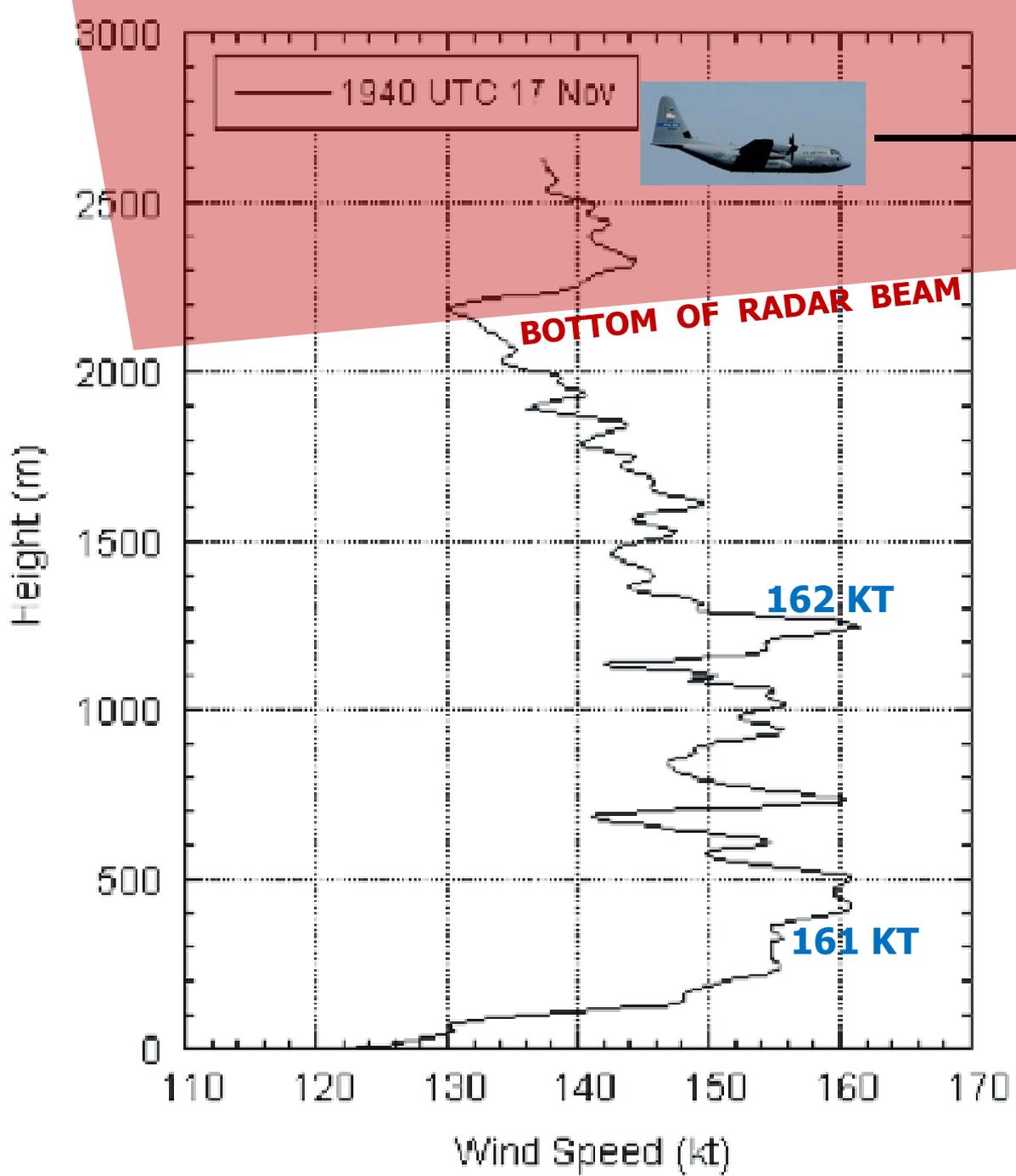
DBZ



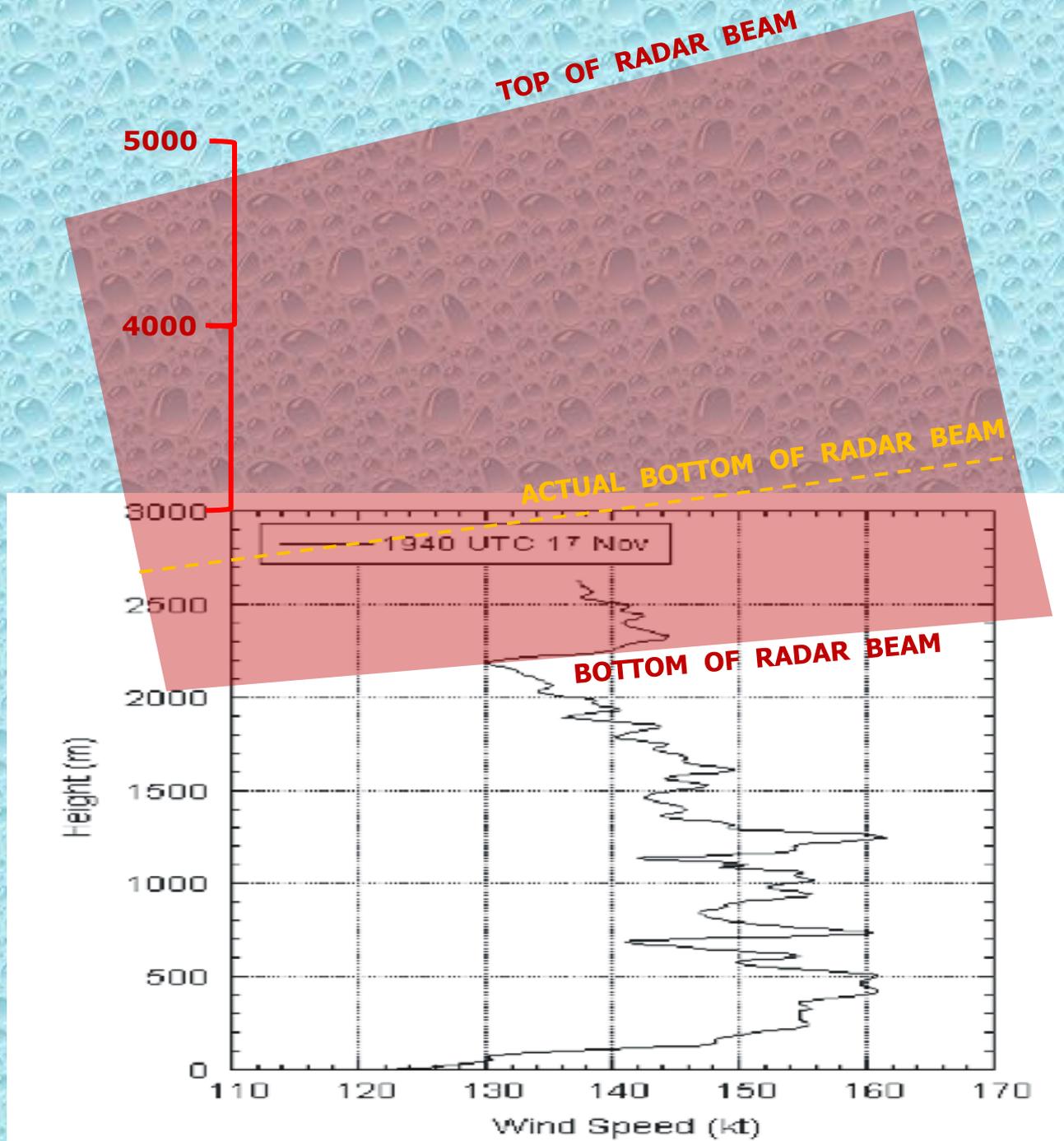
MAX DBZ: 51



Strongest winds were
blowing perpendicular
to the radar beam



Recon flight-level winds: 144 kt

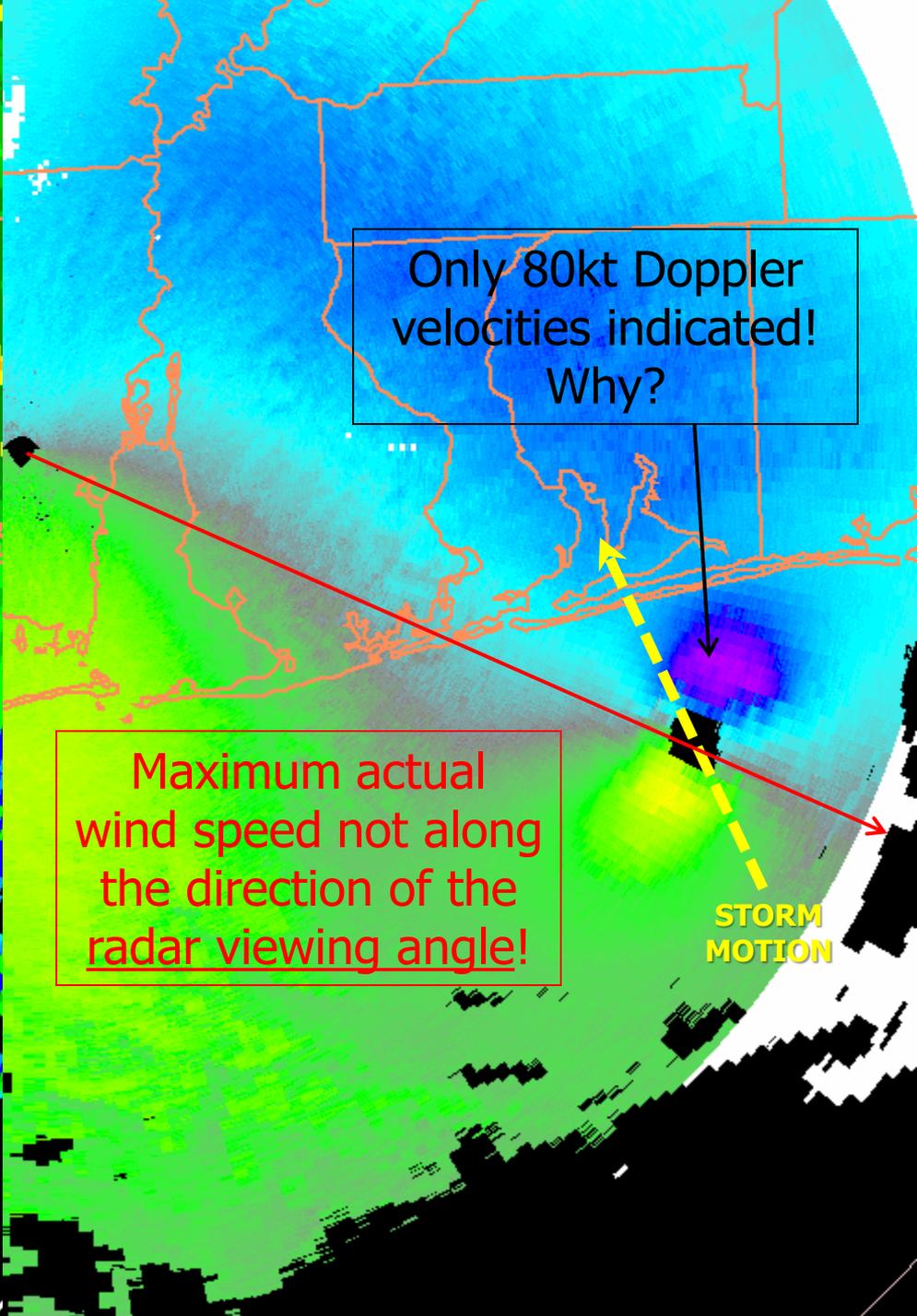
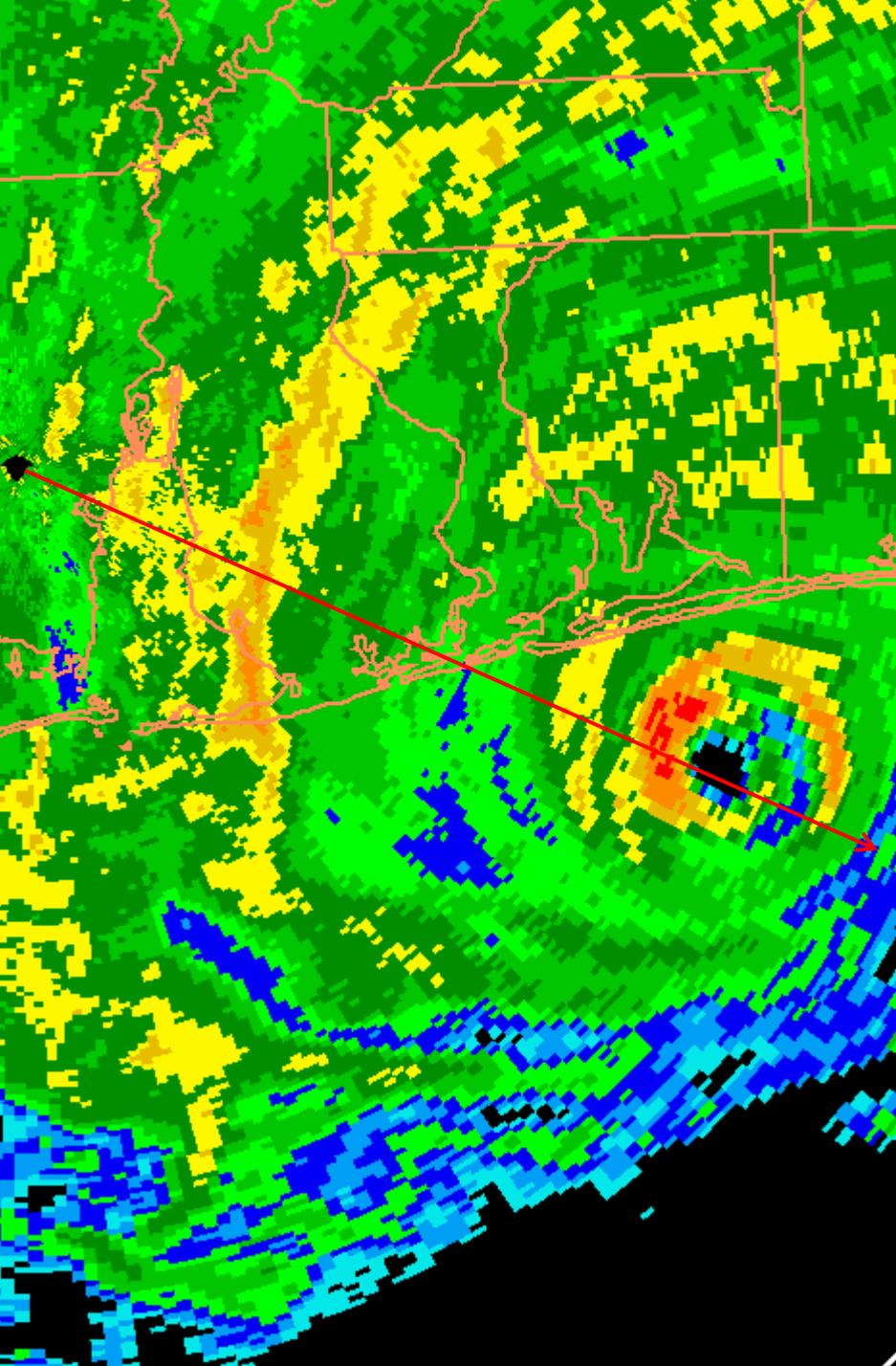


Radar beam is overshooting and not sampling strongest winds below 1,500 meters altitude at 0.5° elevation angle and at sea-level.

However, when San Juan radar antenna altitude of 860 meters is factored in, the bottom of the radar beam is actually at 3,000 meters ASL!

Example -- Hurricane Dennis (2005)

105 kt intensity at landfall in the Florida panhandle as determined by recon aircraft

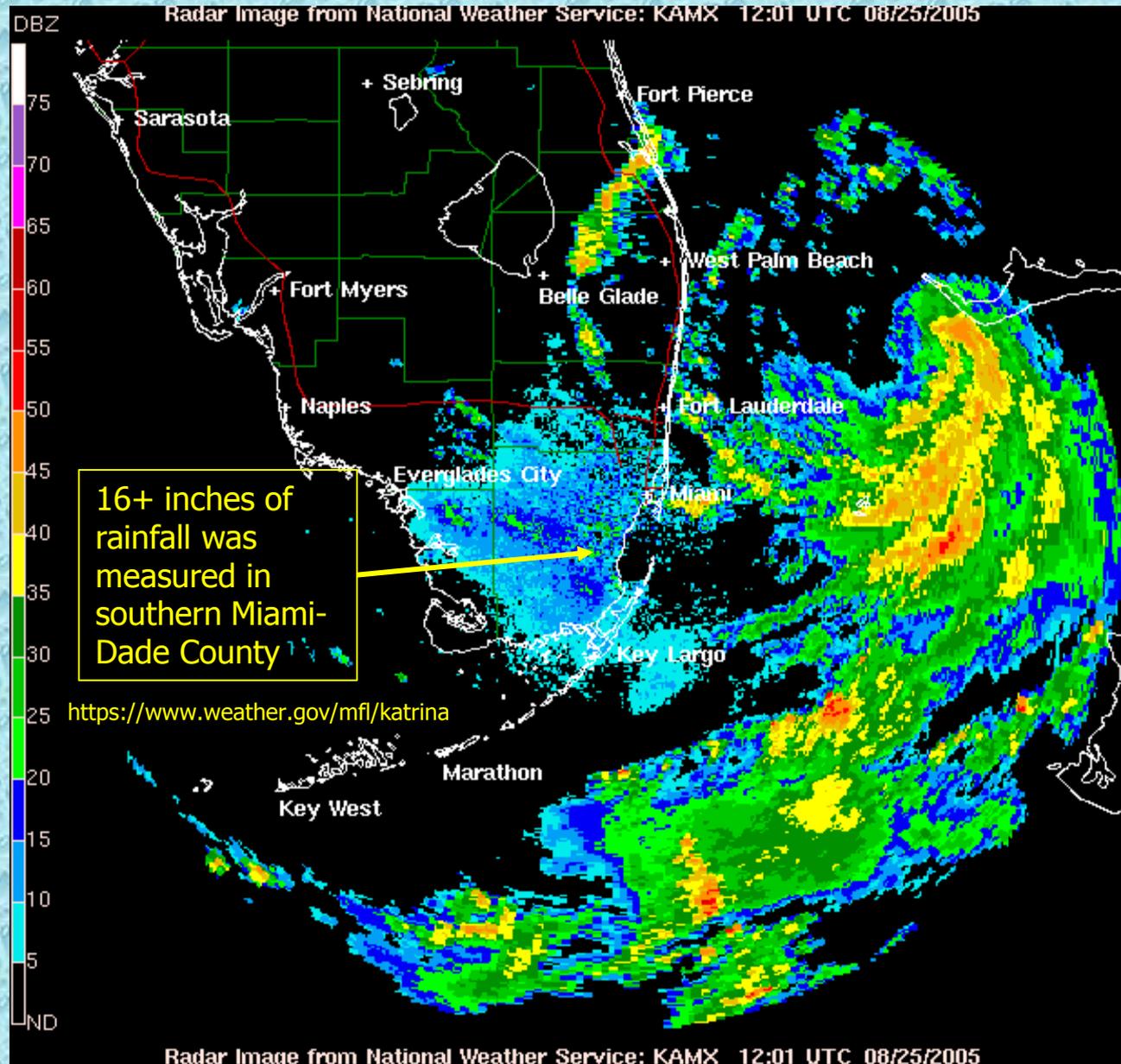


Only 80kt Doppler velocities indicated!
Why?

Maximum actual wind speed not along the direction of the radar viewing angle!

STORM MOTION

Hurricane Katrina (2005) making landfall near Broward & Miami-Dade County line and later moved over the NHC

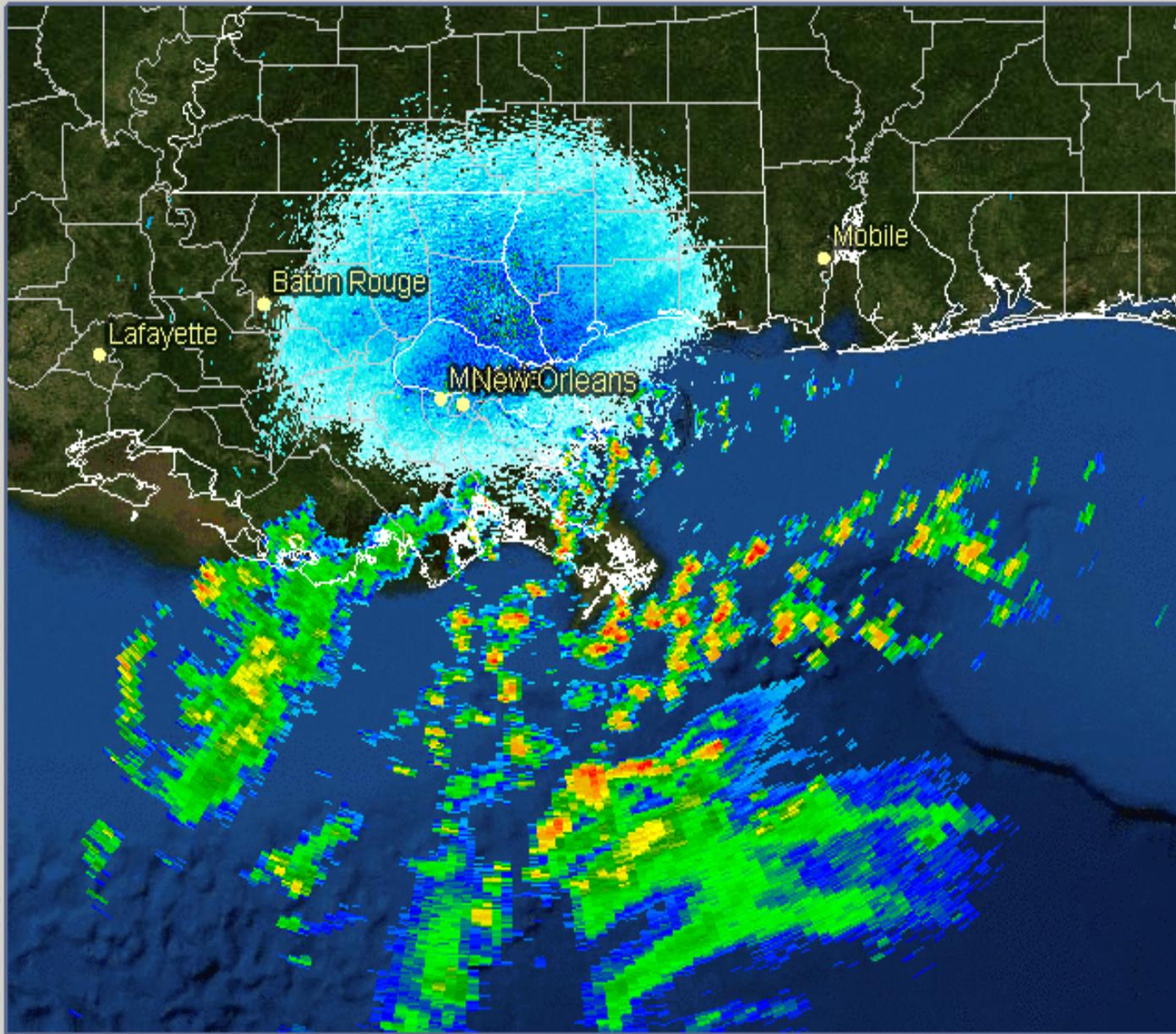


Hurricane Katrina radar observations from NHC – 25 AUG 2005

2005,AUG,25,1700,26.18,79.51,F,CLOSED CIRCULATION WITH DRY SLOT TO N AND W.,RM,KAMX
2005,AUG,25,1728,26.19,79.50,F,MAX WND 65 KT S QUAD AT 4730 FT,RM,KAMX
2005,AUG,25,1801,26.16,79.54,F,MAX WND 77 KT S QUAD AT 5075 FT,RM,KAMX
2005,AUG,25,1829,26.16,79.63,F,WEAK REF NW SEMICIRCLE...MAX WND 75 KT SE QUAD,RM,KAMX
2005,AUG,25,1902,26.14,79.67,F,80 KTS S OF CTR AT 3900 FT,RM,KAMX
2005,AUG,25,1929,26.14,79.74,F,85 KT MAX S QUAD AT 3417 FT JUST S OF CTR,RM,KAMX
2005,AUG,25,2002,26.07,79.84,F,CENTER RAGGED CMA GOOD ZERO ISODOP,CJM,KAMX
2005,AUG,25,2030,26.00,79.90,G,INBOUND MAX OVER SE BROWARD 75 KT 2300',CJM,KAMX
2005,AUG,25,2058,25.98,79.95,G,STRONGEST CONV S EYEWALL KEY BISCAYNE AREA,CJM,KAMX
2005,AUG,25,2134,25.99,79.96,G,CENTRAL CONV BECOMING MORE SYMMETRICAL,CJM,KAMX
2005,AUG,25,2204,25.97,80.03,G,- -,CJM,KAMX
2005,AUG,25,2231,25.96,80.10,G, **CENTER OVER COAST BROWARD-DADE COUNTY** LINE,CJM,KAMX
2005,AUG,25,2304,25.96,80.16,G,HIGHEST WINDS OFFSHORE ABOUT 70 KT 1500',CJM,KAMX
2005,AUG,25,2332,25.89,80.24,G,- -,CJM,KAMX
2005,AUG,26,0000,25.88,80.31,G,- -,CJM,KAMX
2005,AUG,26,0034,25.78,80.39,G,NHC IN EYE - CALM OUTSIDE,CJM,KAMX
2005,AUG,26,0101,25.73,80.46,G,75 KT INBOUND OVER CENTRAL DADE 700 FT ,CJM,KAMX
2005,AUG,26,0126,25.70,80.53,G,- -,CJM, KAMX
2005,AUG,26,0201,25.64,80.60,G,TIGHT VELOCITY COUPLET STILL EVIDENT IN VEL ,CJM,KAMX
2005,AUG,26,0228,25.61,80.71,G,- -,CJM,KAMX
2005,AUG,26,0301,25.58,80.82,G,80 KT AT 900 FT OUTBOUND,CJM,KAMX
2005,AUG,26,0326,25.56,80.93,G,74 KT AT 1600 FT,HDC,KAMX

Example -- Hurricane Katrina (2005)

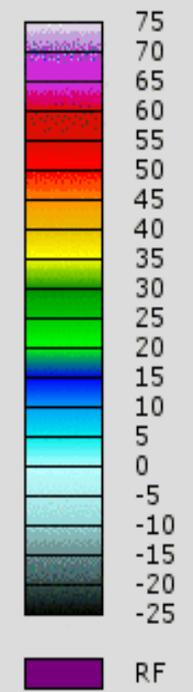
Landfall along the southeast Louisiana coast

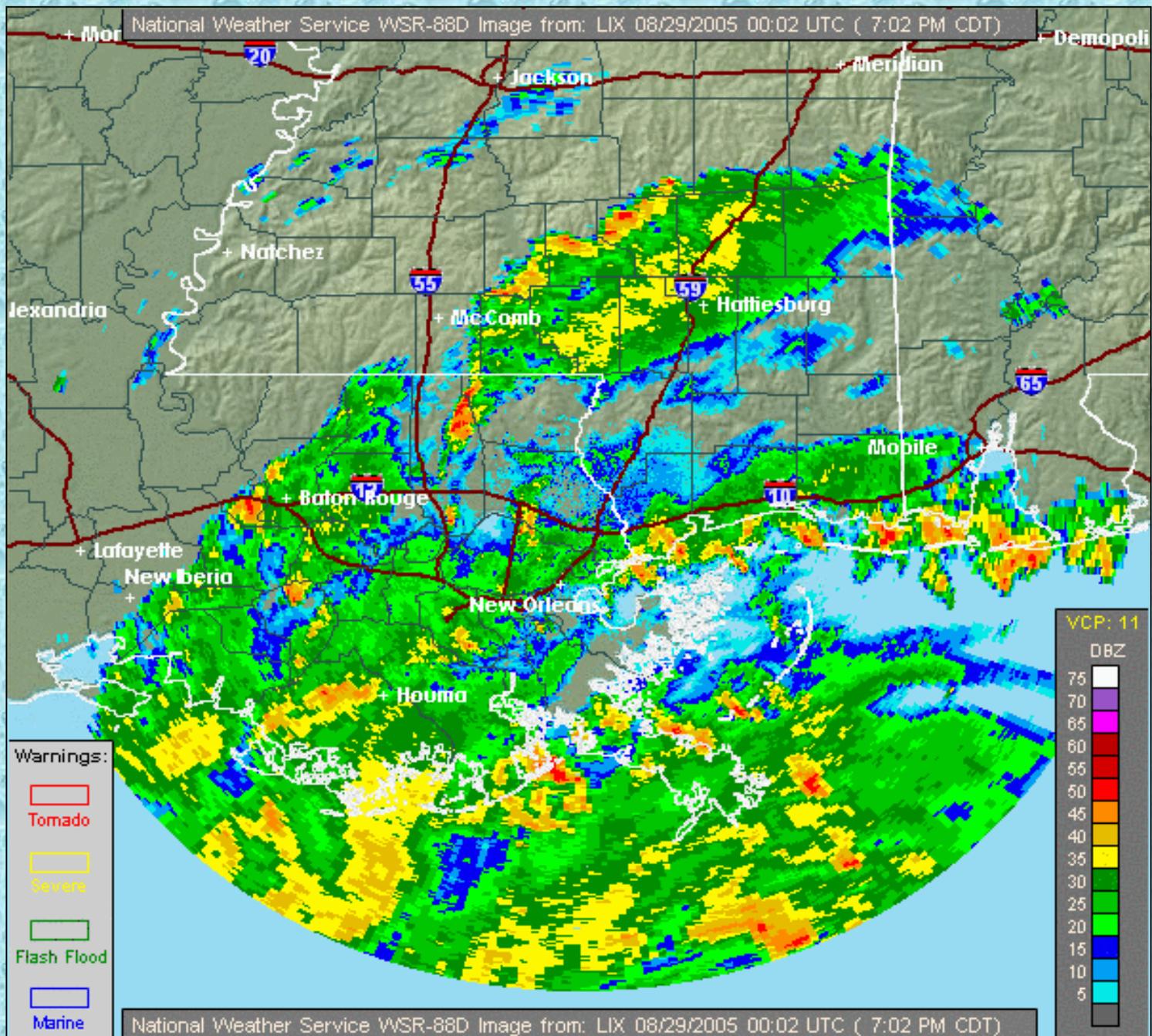


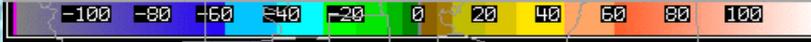
NEXRAD LEVEL-II
KLIX - NEW ORLEANS, LA
08/28/2005 17:04:58 GMT
LAT: 30/20/13 N
LON: 89/49/33 W
ELEV: 24 FT
VCP: 11

REFLECTIVITY
ELEV ANGLE: 0.38

Legend: dBZ

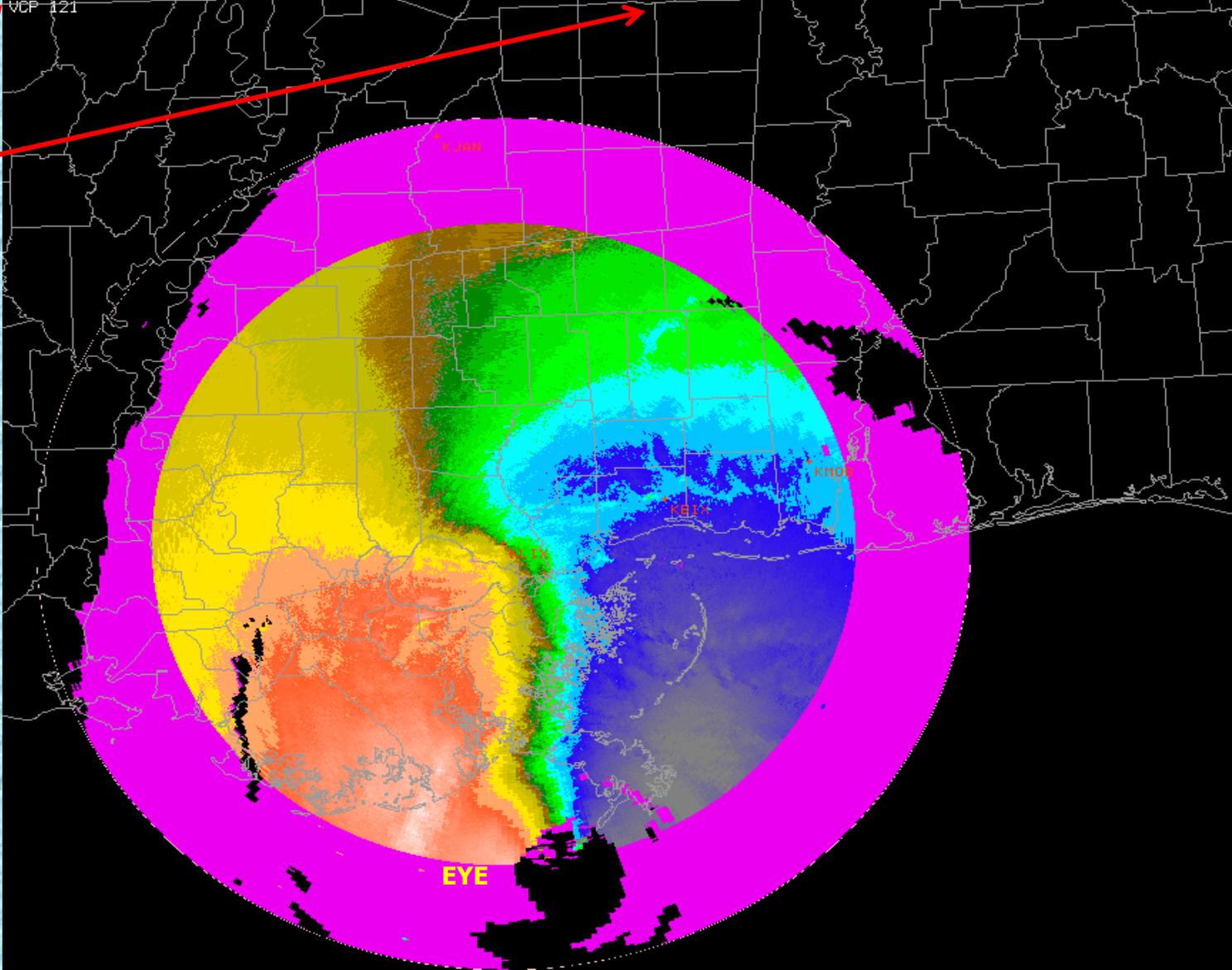


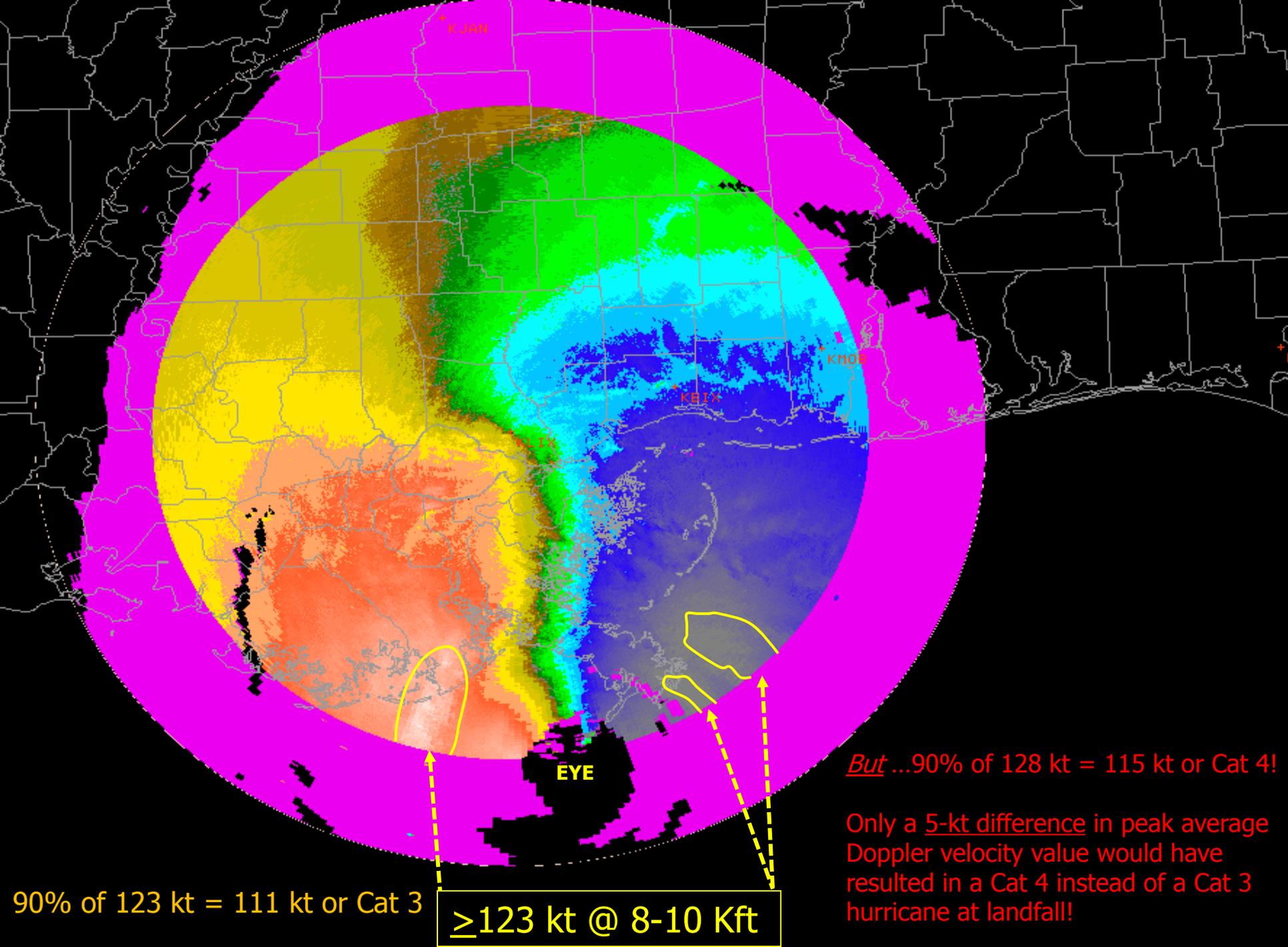




VCP 121

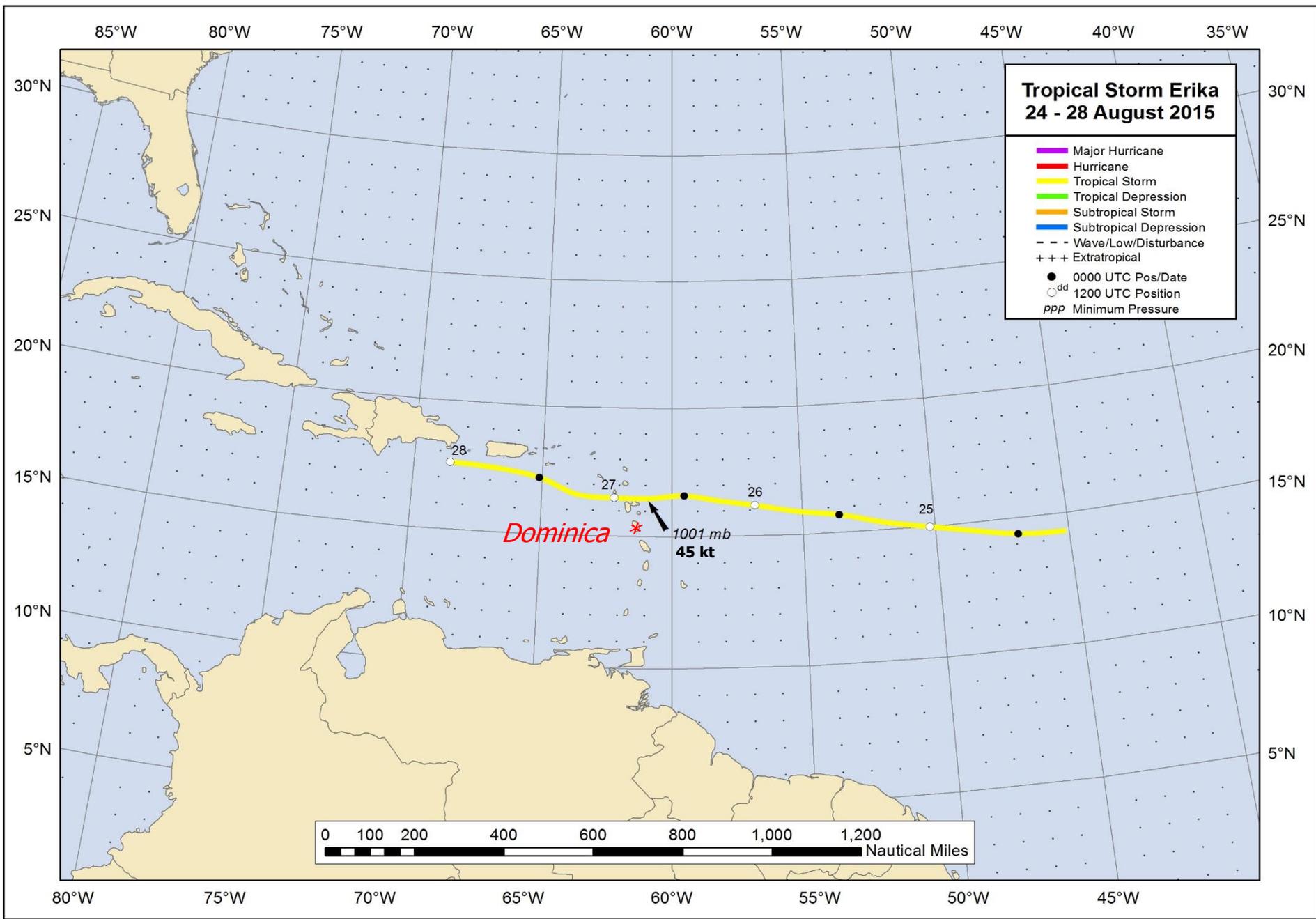
Maximum velocity
Range is -123 kt to
+123 kt due to not
changing velocity
increment to expand
detectable velocity
range to ± 248 kt





Example -- Tropical Storm Erika (Aug 2015): Heavy Rainfall Event

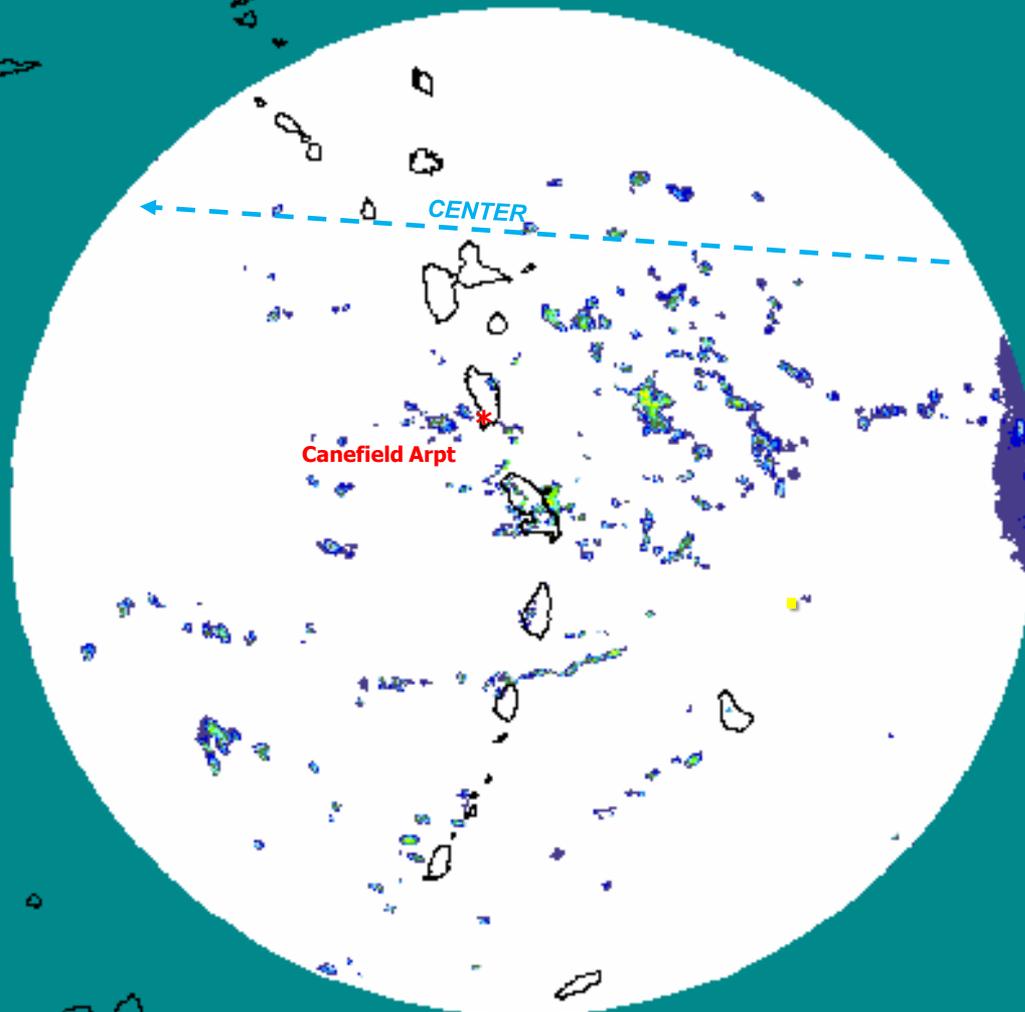
- Cyclone passed through the northern Leeward Islands on 27 August 2015 with only 40-45 kt winds.
- Produced 12.62 inches (320 mm) of rainfall in ~12 hours (0600-1800 UTC) measured at Canefield Airport (TDCF) on southwestern coast of Dominica.
- Caused flash flooding and mudslides on Dominica, damaged or destroyed 271 houses, and caused major damage to roads, bridges and other infrastructure.
- Damage estimated to be to US\$500 million on Dominica.
- 30 people killed and 574 persons left homeless.



Mosaïque radar

Guadeloupe / Martinique

26 / 08 / 2015
14H00 UTC



REFLECTIVITY

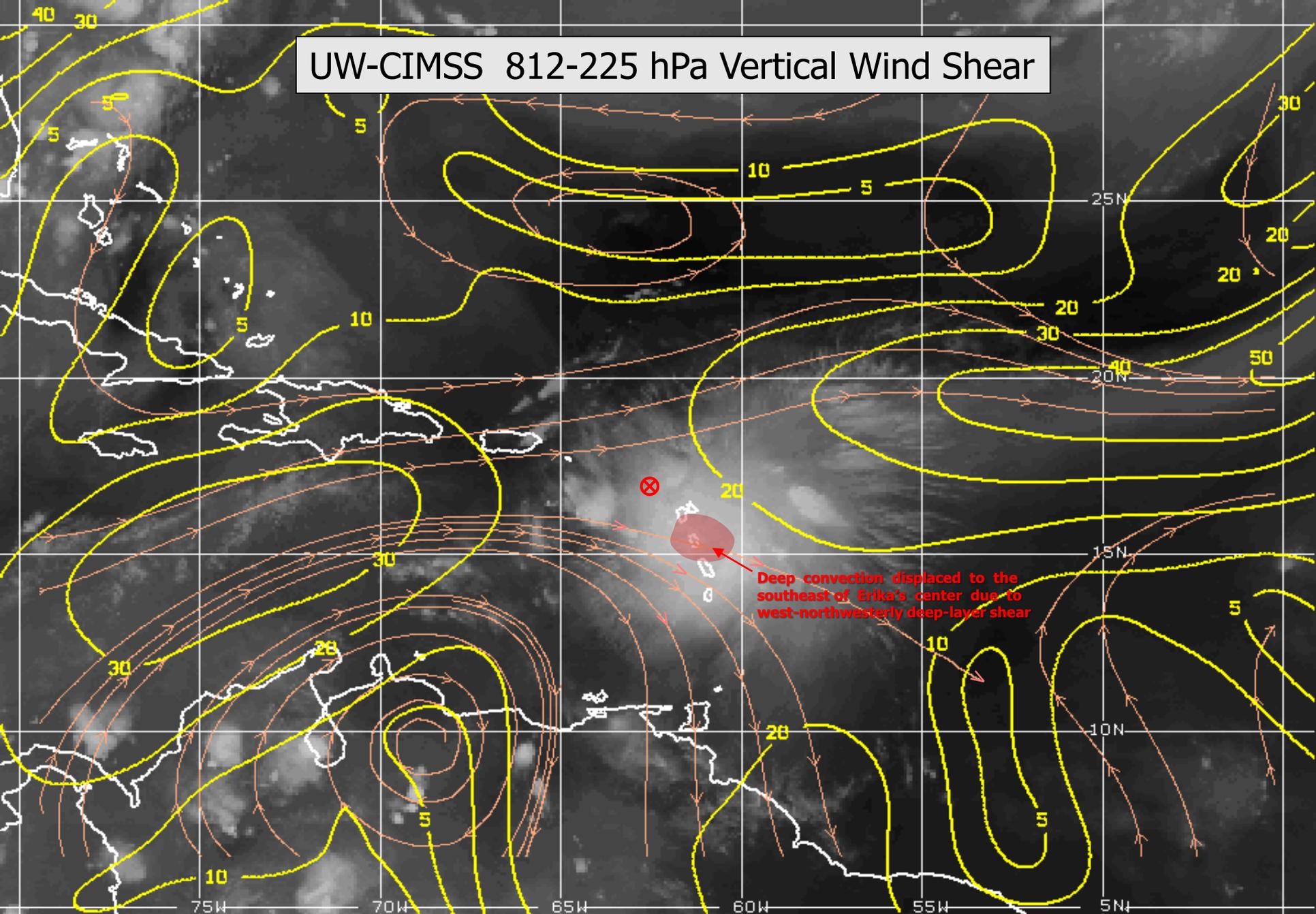
| | |
|---------------|--------------|
| Red | 62.0- 66.0+> |
| Orange-red | 58.5- 62.0 |
| Orange | 55.0- 58.5 |
| Light orange | 51.5- 55.0 |
| Yellow-orange | 48.0- 51.5 |
| Yellow | 44.5- 48.0 |
| Light yellow | 41.0- 44.5 |
| White | 37.5- 41.0 |
| Light green | 34.0- 37.5 |
| Green | 30.5- 34.0 |
| Dark green | 27.0- 30.5 |
| Cyan | 23.5- 27.0 |
| Light blue | 20.0- 23.5 |
| Blue | 16.5- 20.0 |
| Dark blue | 13.0- 16.5 |

dBZ

RANGE : 400 km



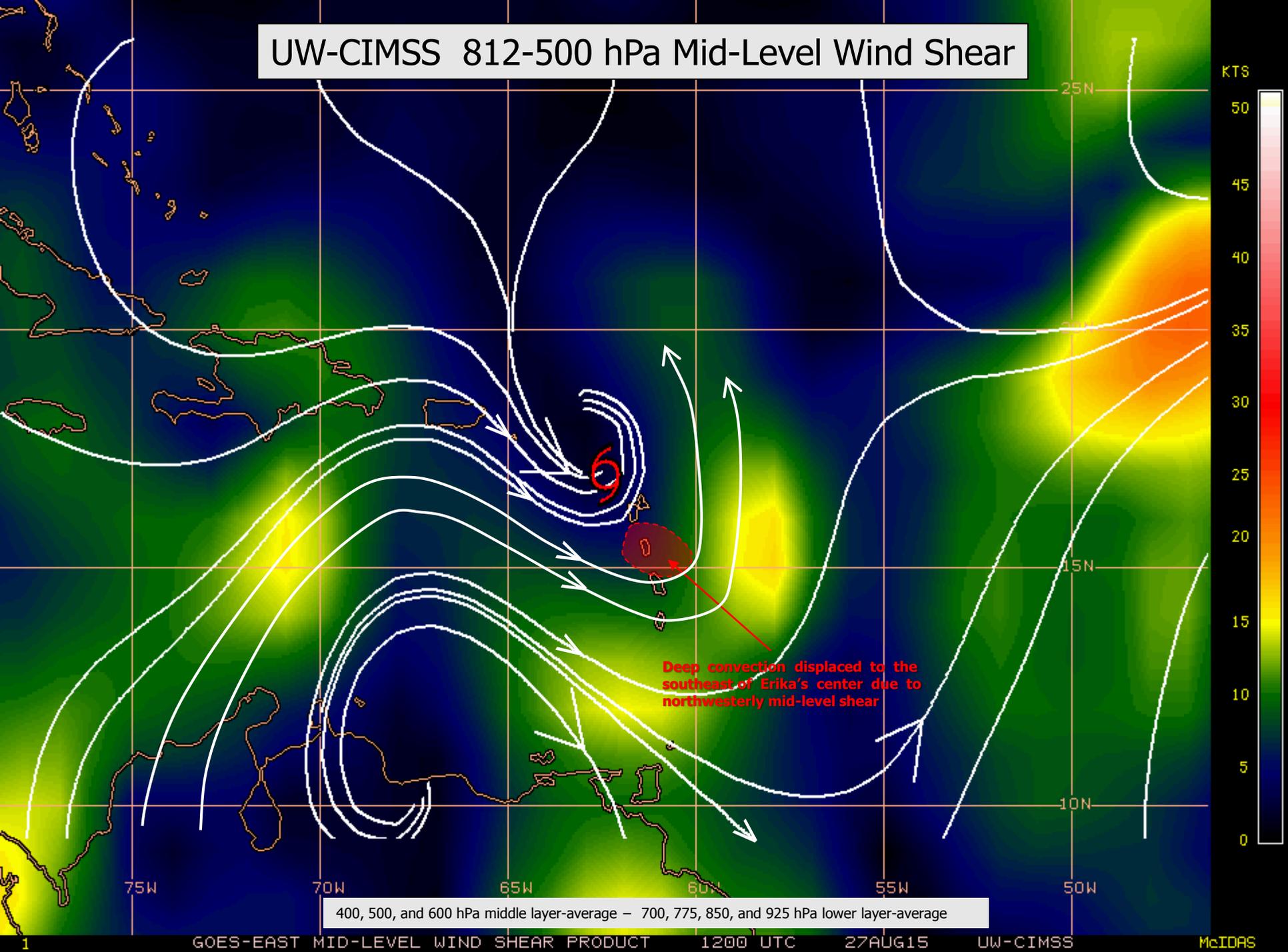
UW-CIMSS 812-225 hPa Vertical Wind Shear

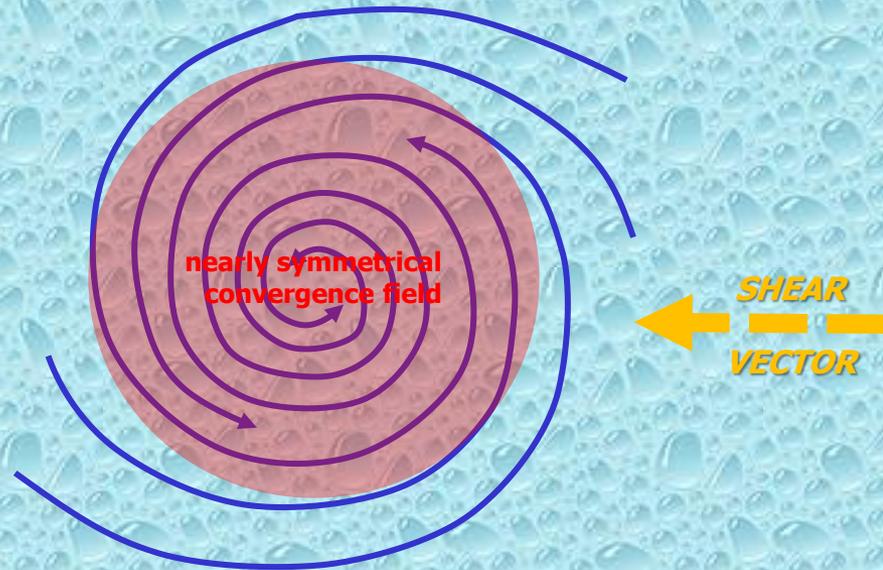


Deep convection displaced to the southeast of Erika's center due to west-northwesterly deep-layer shear

150, 200, 250, and 300 hPa upper layer-average – 700, 775, 850, and 925 hPa lower layer-average

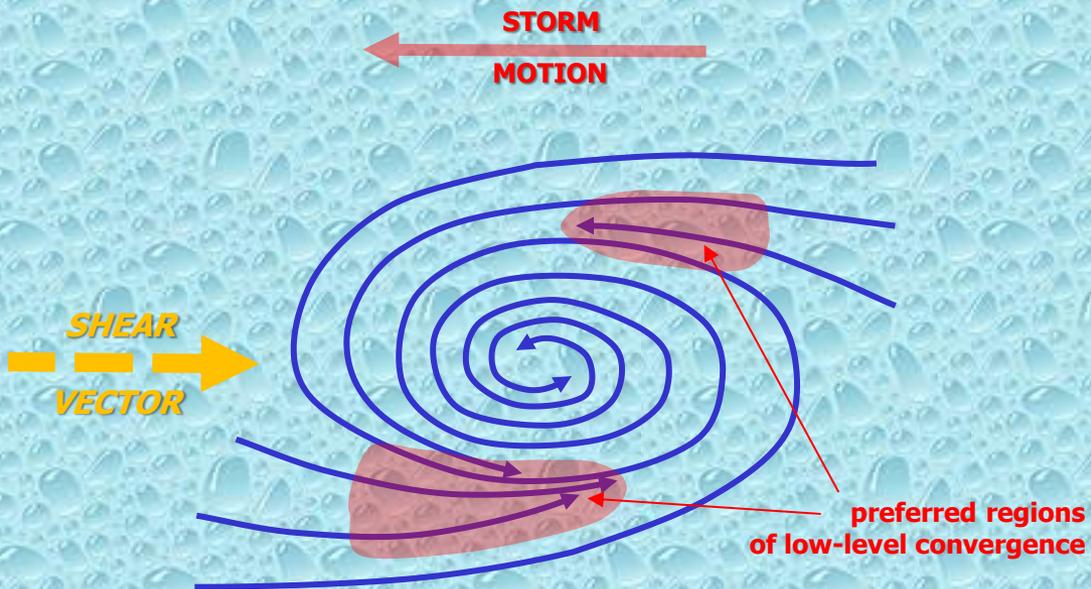
UW-CIMSS 812-500 hPa Mid-Level Wind Shear





Typical symmetrical TC

- Low shear
- Shear vector in same direction as TC motion
- Balanced low-level convergence field



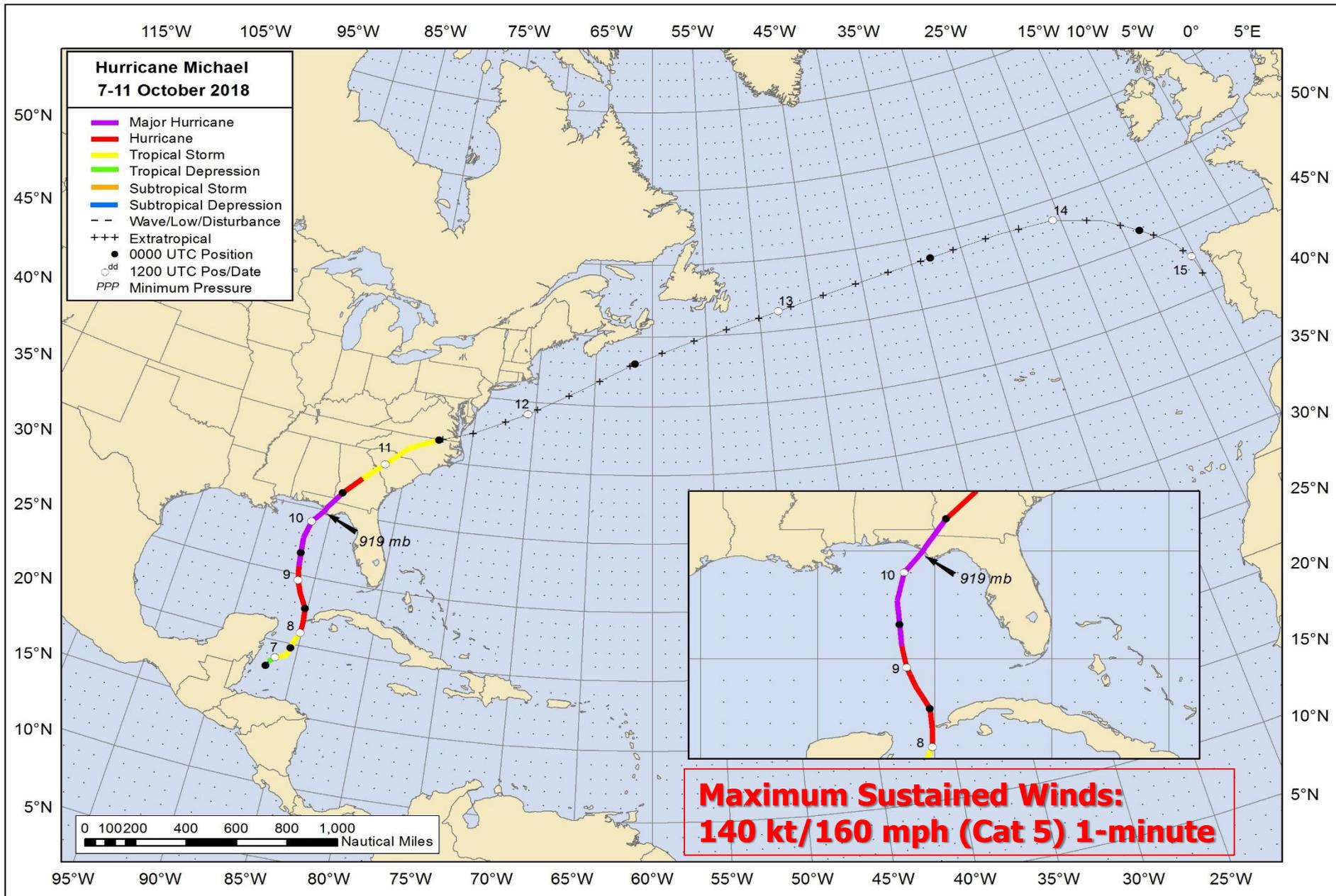
Sheared TC

- Moderate to strong shear
- Shear vector in opposite direction as TC motion
- Causes asymmetry of surface pressure/wind fields
- Creates unbalanced low-level convergence field

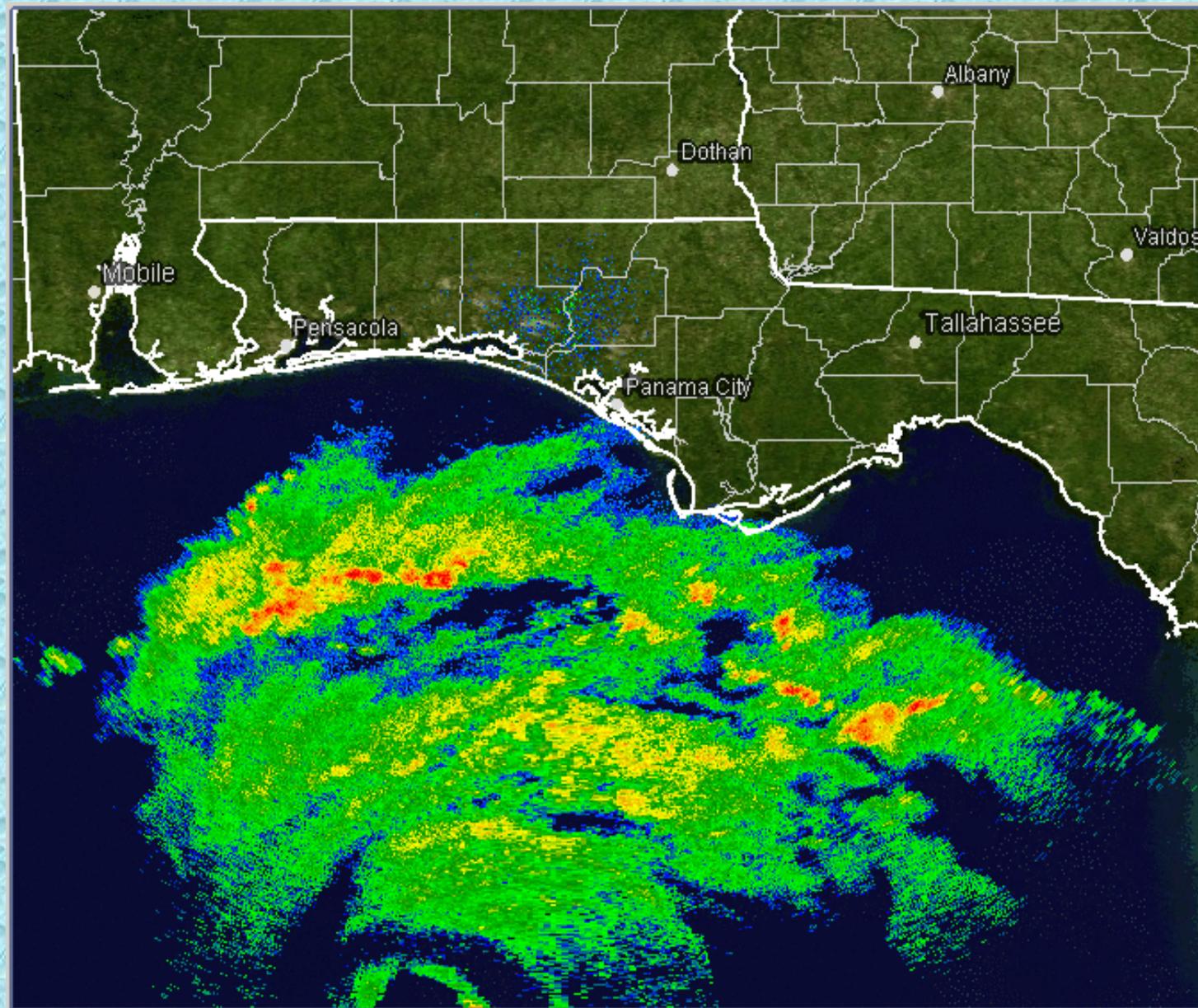
Cause of the heavy rainfall on Dominica

- Localized storm-scale forcing caused by development of low-level convergence zone over/near island.
- Localized mesoscale forcing caused by development of leeside vortex west of and over island.
- Persistent orographic lifting of very moist and unstable air mass caused by low-level westerly winds on south side of TS Erika.
- Strong west-northwesterly 850-200 mb vertical wind shear of 23-33 kt (SHIPS model) displaced convective mass toward the south and east side of Erika's circulation.

Category 5 Hurricane Michael, 7 - 11 October 2018



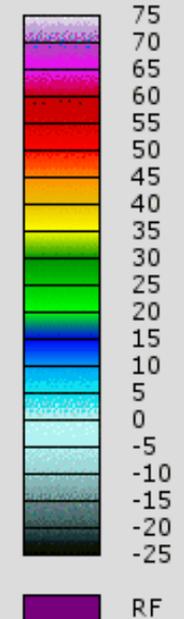
Hurricane Michael, 10 October 2018 (long range)



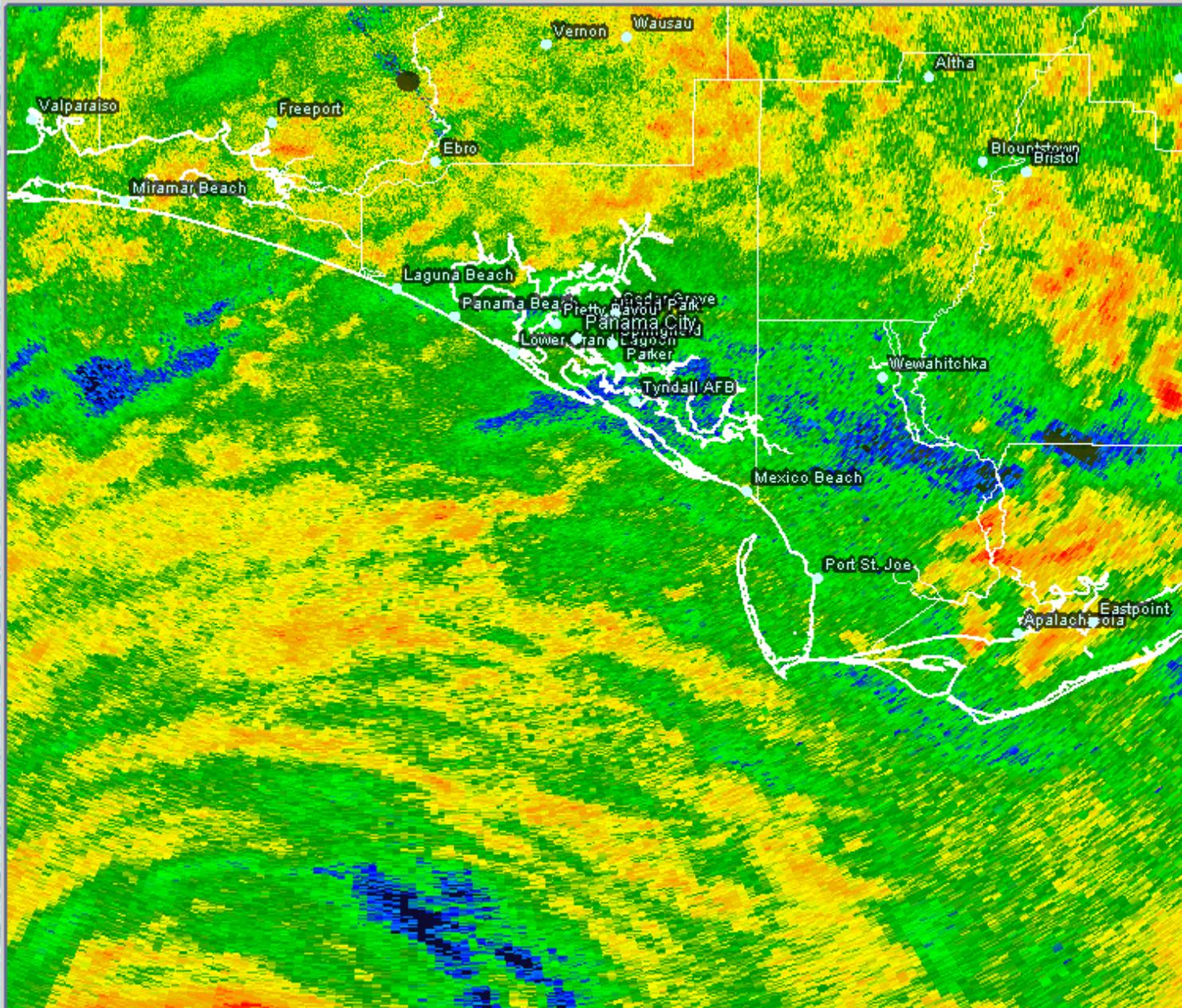
NEXRAD LEVEL-II
KEVX - EGLIN AFB, FL
10/10/2018 04:01:51 GMT
LAT: 30/33/51 N
LON: 85/55/17 W
ELEV: 140 FT
VCP: 212

REFLECTIVITY
ELEV ANGLE: 0.57
SWEEP TIME: 04:01:55 GMT

Legend: dBZ



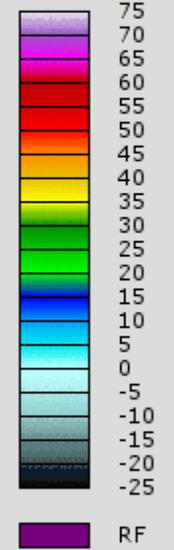
Hurricane Michael, 10 October 2018 (short range)



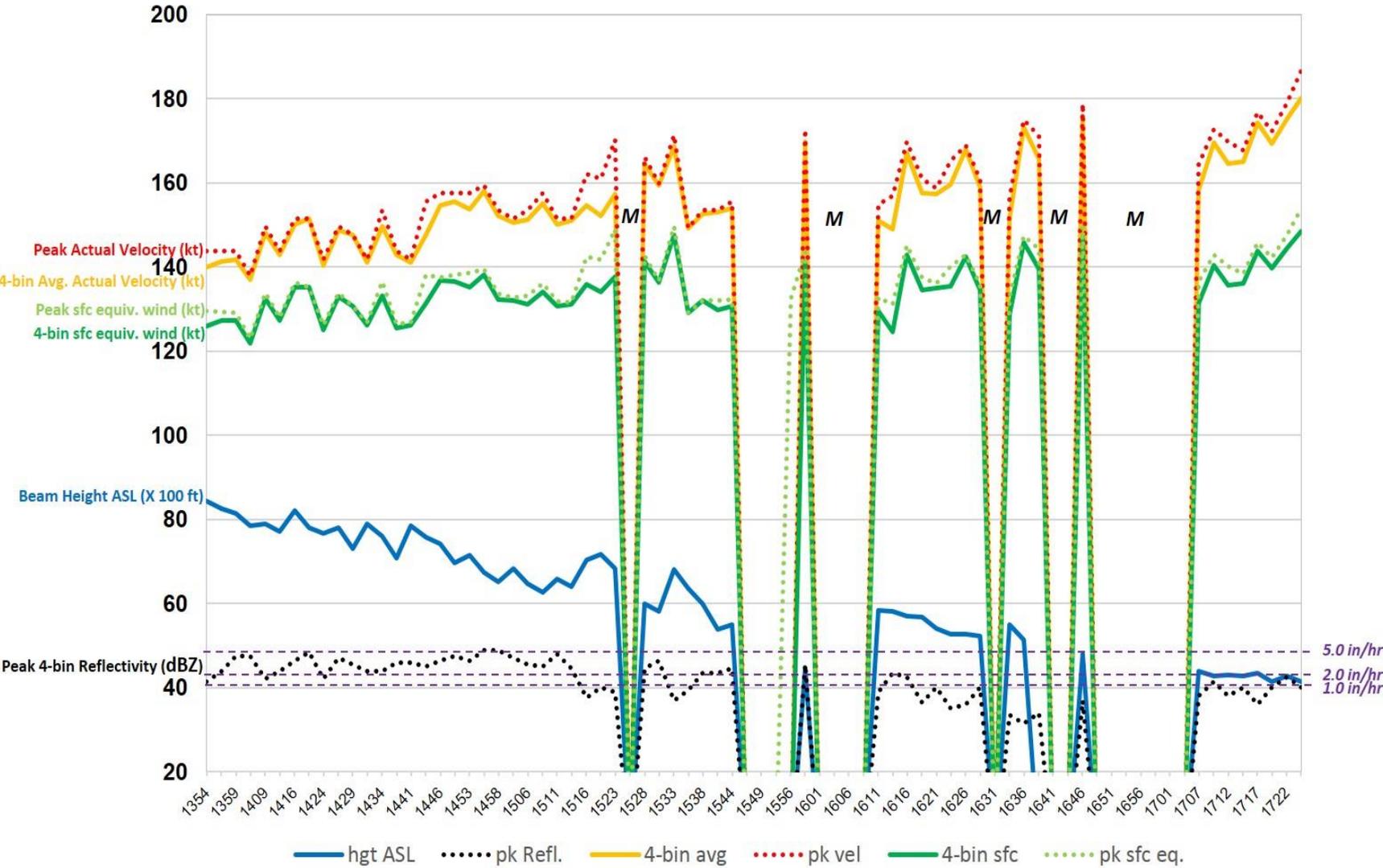
NEXRAD LEVEL-II
KEVX - EGLIN AFB, FL
10/10/2018 12:00:40 GMT
LAT: 30/33/51 N
LON: 85/55/17 W
ELEV: 140 FT
VCP: 212

REFLECTIVITY
ELEV ANGLE: 0.57
SWEEP TIME: 12:00:44 GMT

Legend: dBZ



H. Michael -- KEVX 0.5 deg elev. pre-landfall Doppler radar analysis, 10 OCT 2018 – se quadrant

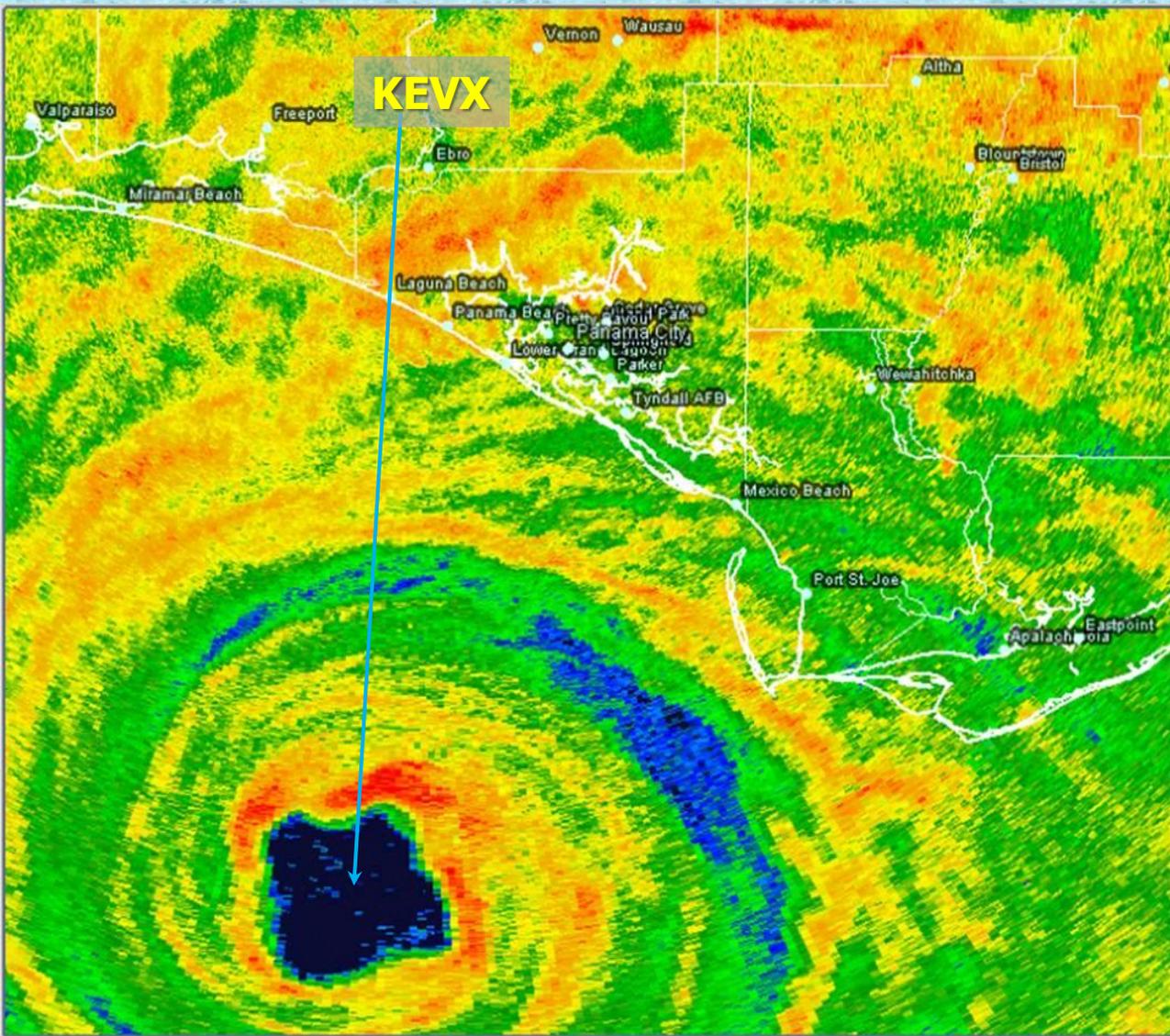


The KEVX WSR-88D Doppler radar data analysis and associated equivalent surface wind speed conversions of the undisturbed tangential wind flow in the southeastern quadrant (090-150° true) leading up to landfall indicate that Hurricane Michael was strengthening right up until landfall occurred at approximately 1730 UTC 10 OCT 2018.

The red-shaded area indicates the time period where original V_{Doppler} values were not converted to V_{actual} values due to AWIPS-II data ingest and display issues; this time period will eventually be converted in the future. However, the wind speeds shown will likely be lower than the converted V_{actual} values.

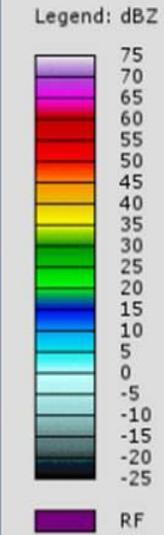
No V_{Doppler} and V_{actual} values were obtained time periods where the tangential wind flow was perturbed by eyewall mesovortices (labeled "M") and, thus, making the values there unrepresentative.

The V_{actual} values over the last ~1 h prior to landfall suggest that Michael had sustained surface wind speeds of at least **140 kt**.



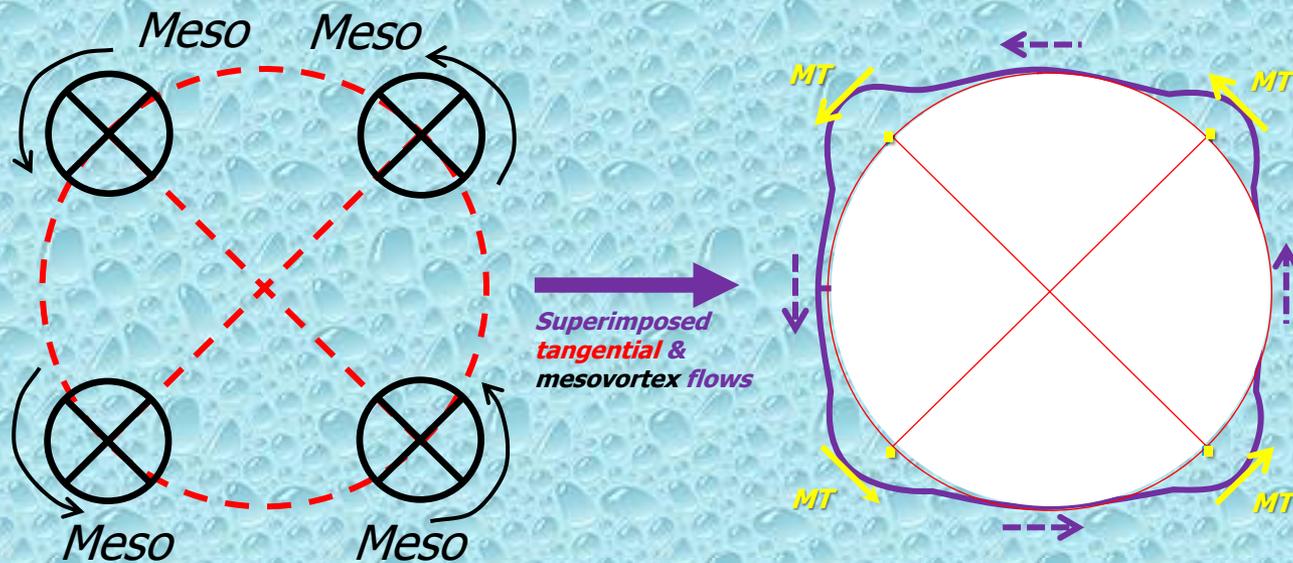
NEXRAD LEVEL-II
 KEVX - EGLIN AFB, FL
 10/10/2018 14:34:04 GMT
 LAT: 30/33/51 N
 LON: 85/55/17 W
 ELEV: 140 FT
 VCP: 212

REFLECTIVITY
 ELEV ANGLE: 0.57
 SWEEP TIME: 14:34:08 GMT



Real-Time example of the combined tangential and four mesovortex flows (MT) associated with Hurricane Michael at 1434:04 UTC 10 OCT.

Some of the MT flows indicated V_{actual} values of 180-200 kt, which corresponds to an equivalent surface wind speed of 153-165 kt using recon adjustment values ranging from 0.825 to 0.850 for the corresponding altitudes of the $V_{Doppler}$ radar bins.



The tangential & mesovortex combined flows can only be accurately assessed at locations **MT** where both flows directions exactly coincide, thus allowing for symmetrical/circular flow to be assumed at those points.

This allows for an accurate assessment of the Cosine of the Radar Viewing Angle (RVA) and, therefore, V_{actual} to be calculated.

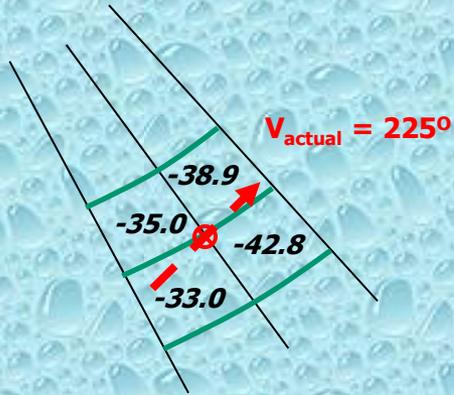
Eyewall mesovortex wind speed data *were not included* in the computation of V_{actual} values.

However, there were at least 5 cases where the wind direction of the tangential winds and the mesovortex winds coincided, allowing for the calculation of peak combined flows and V_{actual} values, which ranged from ~ 180 - 200 kt, resulting in equivalent surface wind speed estimates of 153-165 kt.

Some consideration and weight should be given to the equivalent surface wind speeds associated with the eyewall mesovortices (eddy flow) since the temporal and spatial scales of those features were very similar to scale of the undisturbed eyewall tangential flow between the mesovortices.

Recon Wind Data vs. WSR-88D (V_{Doppler}) & V_{actual} Direction

AF301 1514A MICHAEL HDOB 27 20181010
 171930 3000N 08534W 6970 02432 9177 +191 +097 194027 030 053 002 03
 172000 2959N 08532W 6960 02446 9181 +191 +100 193030 031 /// /// 03
 172030 2958N 08532W 6968 02435 9173 +197 +105 208032 033 049 002 00
 172100 2956N 08531W 6963 02446 9171 +205 +121 229049 065 083 001 03
 172130 2955N 08530W 6981 02466 9242 +168 +135 238107 122 099 001 00
 172200 2954N 08528W 6967 02555 9327 +140 +139 231142 149 121 002 05
 172230 2953N 08527W 6973 02621 9437 +137 +136 225150 152 123 005 03
 172300 2952N 08525W 6977 02684 9527 +124 //// 218141 148 132 007 05
 172330 2951N 08524W 6971 02747 //// +114 //// 219140 146 133 006 05
 172400 2951N 08522W 6976 02789 9613 +128 +118 217132 136 101 002 03
 172430 2950N 08521W 6971 02819 9641 +130 +106 219124 128 092 001 00
 172500 2949N 08522W 6971 02829 9654 +126 +109 222122 123 092 001 00
 172530 2948N 08522W 6967 02844 9667 +125 +109 224119 120 091 001 00
 172600 2947N 08522W 6970 02853 9679 +123 +114 226116 118 088 003 00
 172630 2946N 08522W 6968 02866 9699 +119 +117 229112 115 085 006 00
 172700 2945N 08523W 6967 02875 9719 +118 +118 233108 110 085 007 00



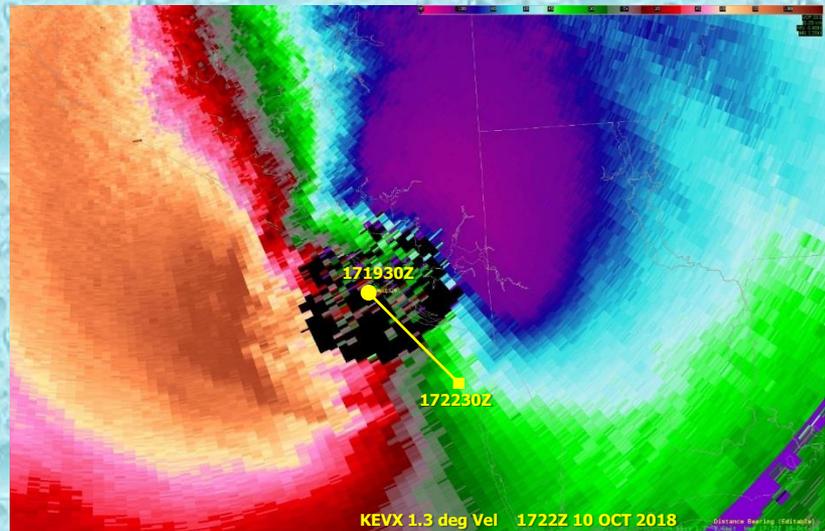
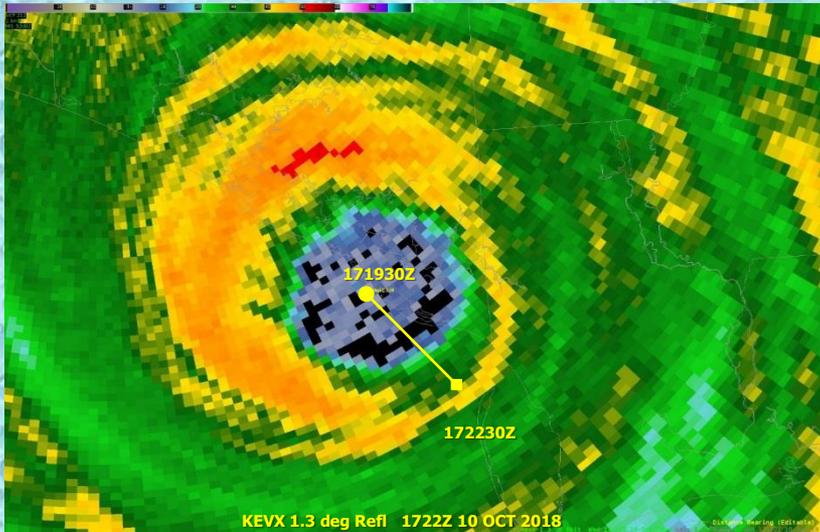
Radar beam height = 8337 ft ASL
 Aircraft altitude/height = 8599 ft ASL

Recon actual wind direction = 225°
 Radar radial = 149°
 Radar viewing angle = 76°
 Cosine $76^\circ = 0.2419$

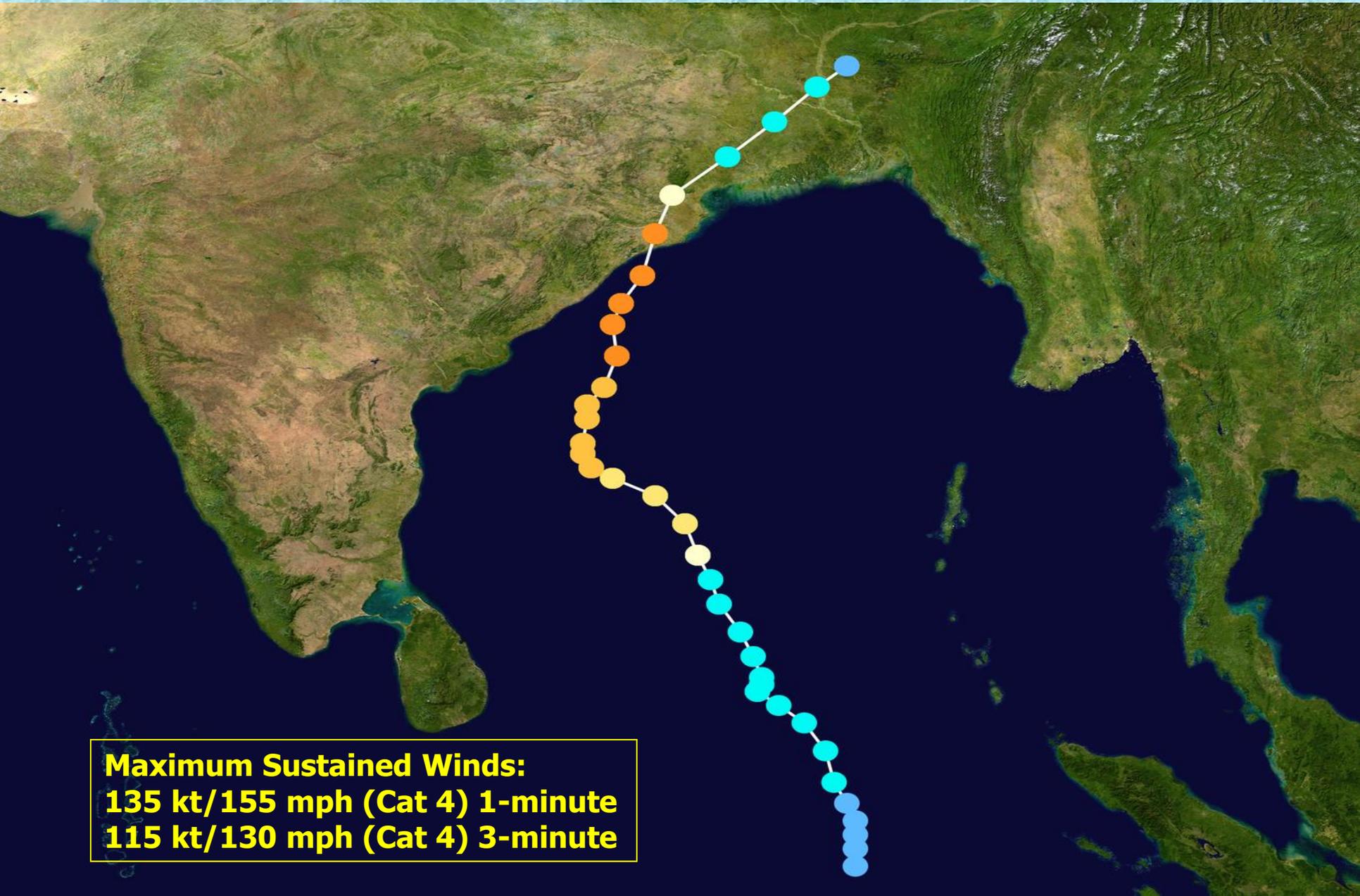
$$V_{actual} = V_{Doppler} / \text{Cosine of angle}$$

$V_{Doppler} 33.0 \text{ kt} \Rightarrow V_{actual} = 136.0 \text{ kt}$
 $V_{Doppler} 35.0 \text{ kt} \Rightarrow V_{actual} = 144.7 \text{ kt}$
 $V_{Doppler} 38.9 \text{ kt} \Rightarrow V_{actual} = 160.8 \text{ kt}$
 $V_{Doppler} 42.8 \text{ kt} \Rightarrow V_{actual} = 176.9 \text{ kt}$

4-bin V_{actual} average = **154.6 kt**
 Recon V_{actual} = **152.0 kt**

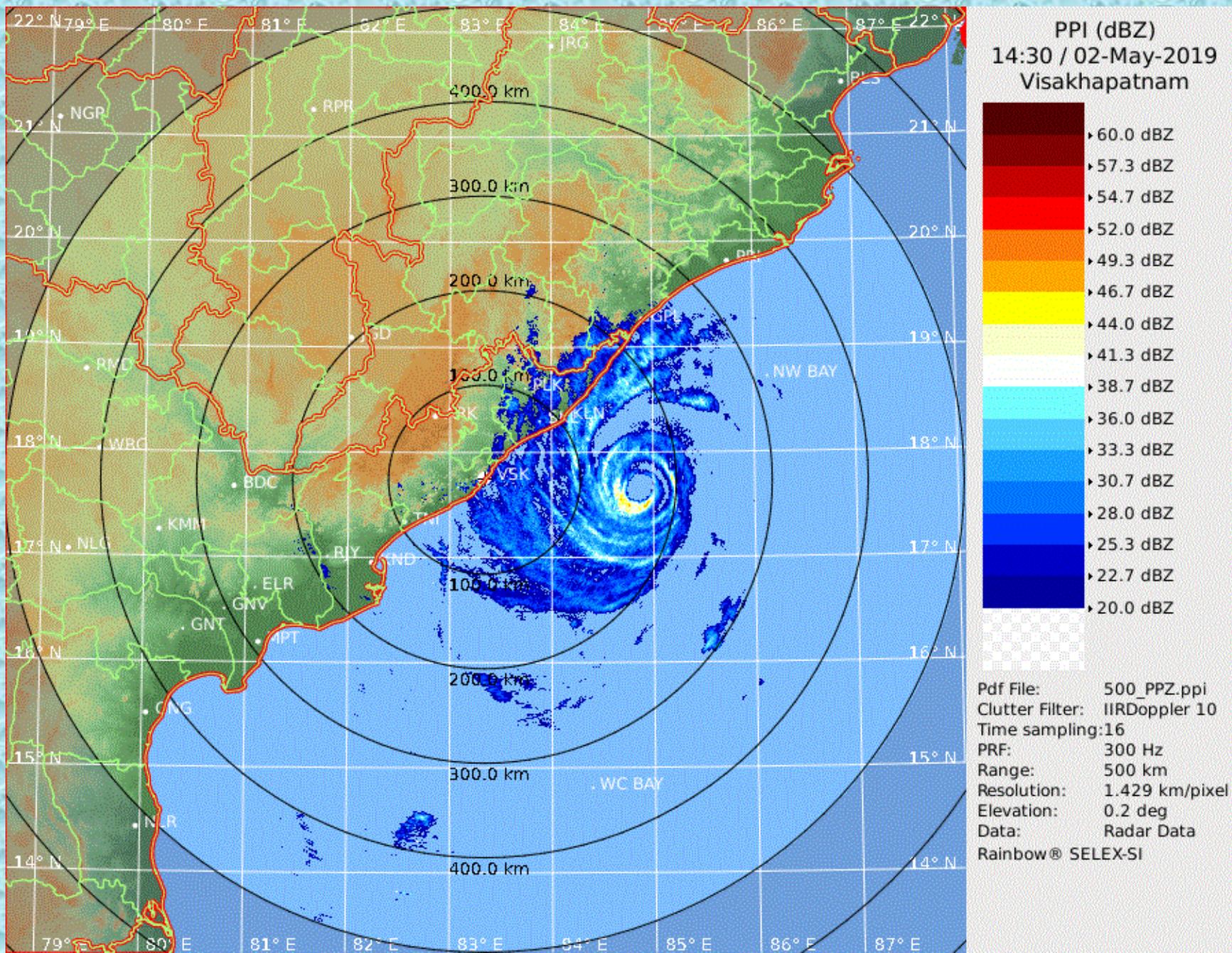


Extremely Severe Tropical Cyclone Fani, 26 April – 5 May 2019



Maximum Sustained Winds:
135 kt/155 mph (Cat 4) 1-minute
115 kt/130 mph (Cat 4) 3-minute

Tropical Cyclone Fani, 02-03 May 2019 (long range)



Tropical Cyclone Fani, 02-03 May 2019 (short range)

gopalpur_weather_2019_05_02_19_31_26_dsfindb.dwr

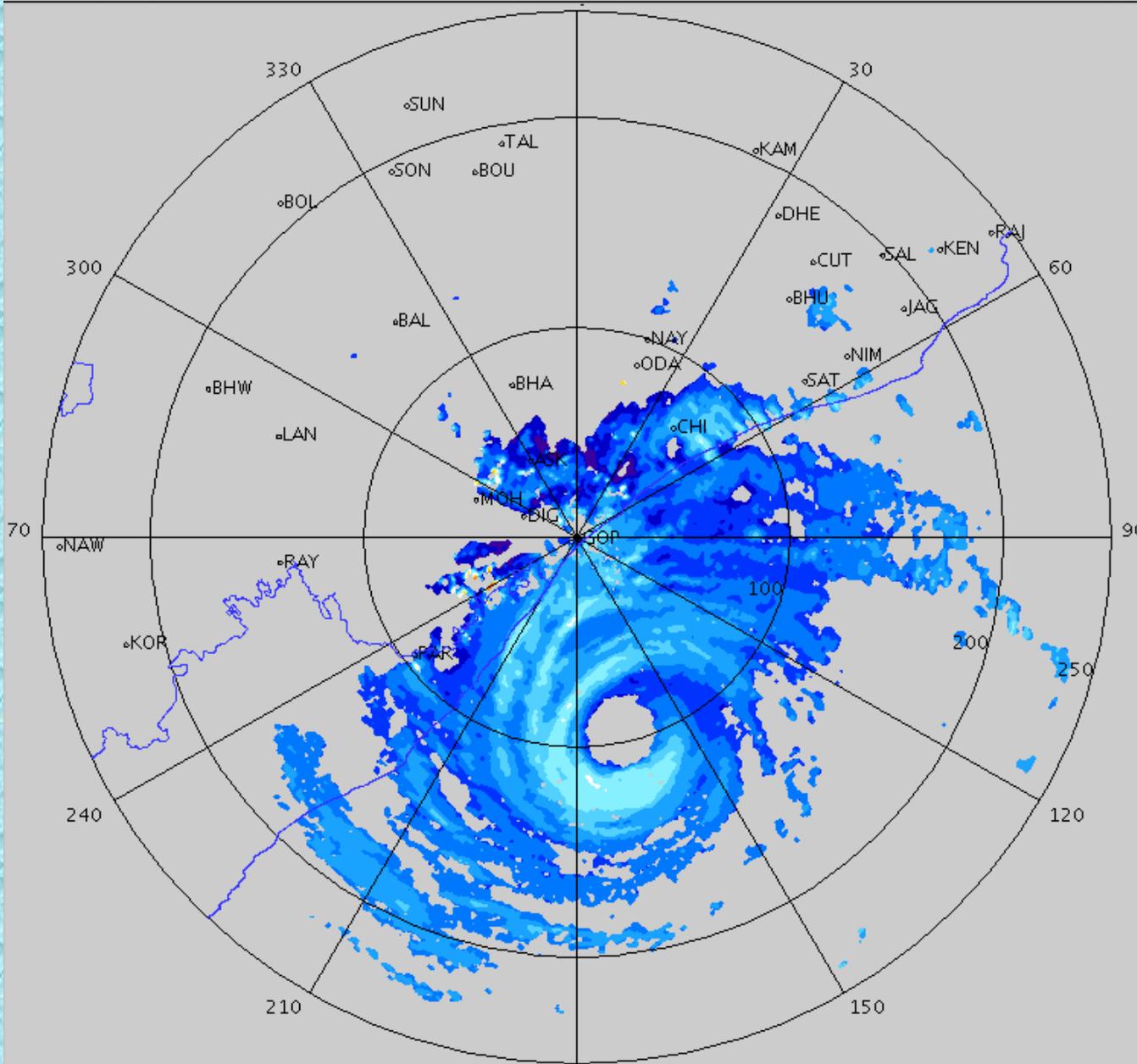
2019/05/02

19:31:26 (UTC)

PPI(Z)

DWR GOPALPUR

(19.2743N , 84.8819E , 26.0000 mts)



| | |
|--------------------|-------------------|
| Elevation | : 0.2 Deg |
| Z (Lo, Hi) | : -31.5, 90.0 dBZ |
| Display Range | : 250 Km |
| Display Res | : 0.7 Km/Pix |
| DDR Enable | : Yes |
| Scan Elev (Lo, Hi) | : 0.2, 21.0 Deg |
| PRF (Lo, Hi) | : 450, 600 Hz |
| Scan Res | : 150 mts |
| Scan RPM | : 2.0 |
| Log Threshold | : 3 dB |
| DTP | : 42 |
| Pulse Width | : 1.0 micro sec |
| Clutter Filter | : IIR-HP(54.0 Hz) |
| SQI | : 0.25 |
| CSR | : 20 dB |
| Scan Range | : 240 Km |
| Preprocessings | : NONE |
| Filters | : NONE |

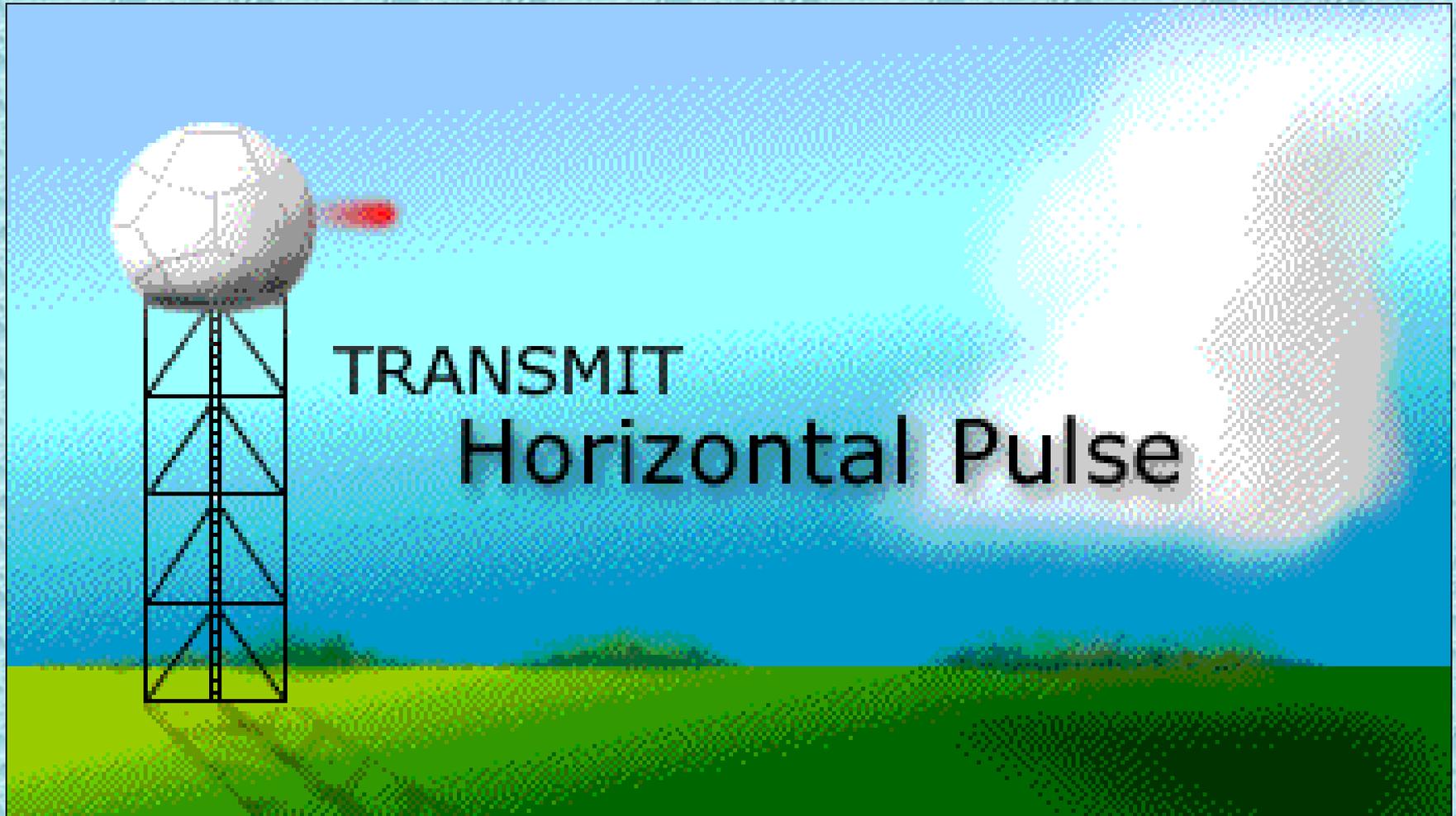
dBZ

| | |
|--|---------------|
| | No Data |
| | 60.67: >65.00 |
| | 56.33: 60.67 |
| | 52.00: 56.33 |
| | 47.67: 52.00 |
| | 43.33: 47.67 |
| | 39.00: 43.33 |
| | 34.67: 39.00 |
| | 30.33: 34.67 |
| | 26.00: 30.33 |
| | 21.67: 26.00 |
| | 17.33: 21.67 |
| | 13.00: 17.33 |
| | 8.67: 13.00 |
| | 4.33: 8.67 |
| | 0.00: 4.33 |



Dual-Polarization Doppler Weather Radars

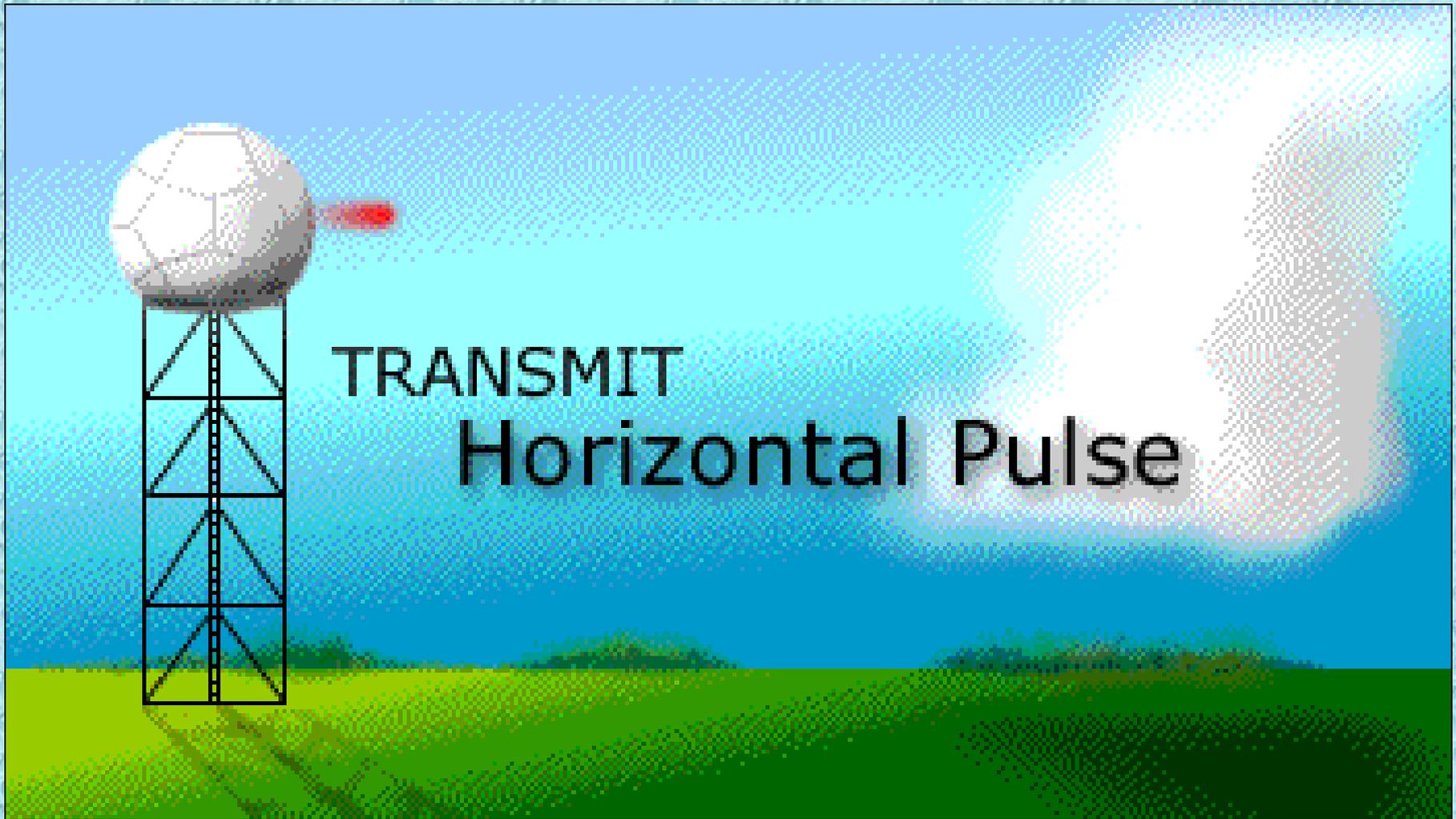
Conventional Horizontal-Polarized Radar



TRANSMIT
Horizontal Pulse

Typical rain drop is a 'flat', oblate spheroid with a wider horizontal axis that returns more energy in the horizontal plane

Dual-Polarized Radar



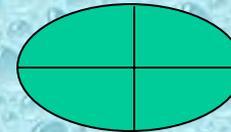
Provides better Z-R relationship and rainfall estimates by determining *precipitation type*

New parameter called "Differential Reflectivity" or " Z_{DR} " helps to determine precipitation type

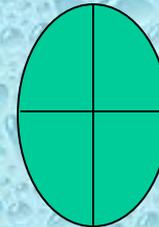
Z_h = horizontal polarized reflectivity

Z_v = vertical polarized reflectivity

$Z_h > Z_v$ for raindrops



$Z_h < Z_v$ for large wet hailstones



$$dBZ_{dr} = 10 \times \log (Z_h / Z_v)$$

Z_{DR} values for meteorological echoes typically range between -2 dB and 6 dB

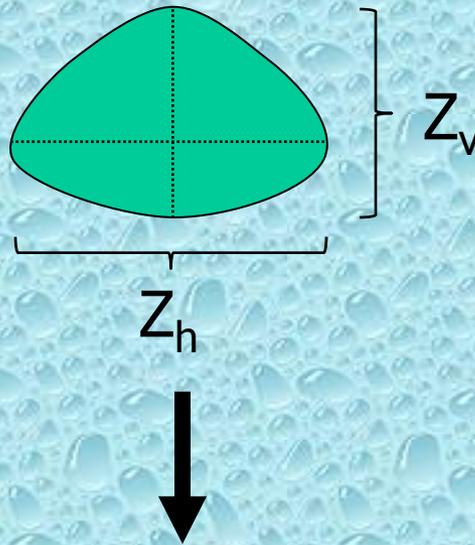
- Values of Z_{DR} well above zero indicate the hydrometeors in the volume are horizontally oriented (e.g., rain) -- meaning their horizontal axis is longer than their vertical axis ($P_h > P_v$).
- Values of Z_{DR} well below zero indicate the hydrometeors in the volume are vertically oriented (e.g., large hail) -- meaning their vertical axis is longer than their horizontal axis ($P_h < P_v$).
- Values of Z_{DR} near zero indicate the hydrometeors in the volume have a nearly spherical shape (e.g., snow, giant hail), in the mean ($P_h \sim P_v$).

Example: Consider a field of large, falling raindrops. The drops tend to fall with an oblate, horizontal orientation. The field of drops, as a whole, will have a larger cross-section of water in the horizontal plane compared to the vertical.

A horizontally-polarized radar pulse will, therefore, backscatter more energy/power in this field of drops than a vertically-polarized pulse will, resulting in more radar return for the horizontal pulse than the vertical pulse.

In this case, $\text{Power}_{\text{hor}} > \text{Power}_{\text{vert}} \Rightarrow Z_h > Z_v$, thus $Z_{\text{DR}} > 0$.

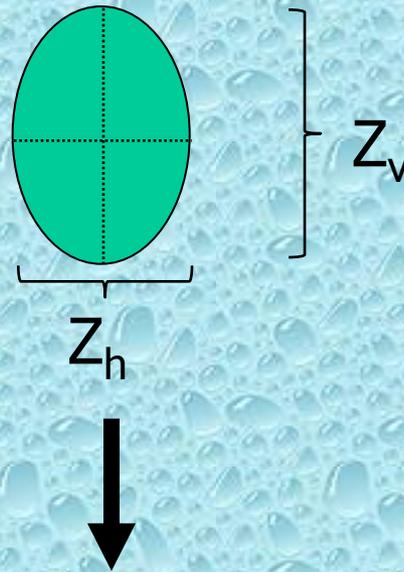
Large Raindrops



Example: $Z_h = 317,000$ (i.e. 55 dBZ) and $Z_v = 100,000$ (i.e. 50 dBZ)

$$\begin{aligned}\text{Therefore, } Z_{DR} &= 10 \log (Z_h/Z_v) \\ &= 10 \log (317000/100000) \\ &= 10 \log (3.17) \\ &= 10 \times 0.501 \\ Z_{DR} &= +5.01\end{aligned}$$

Large Wet Hailstones

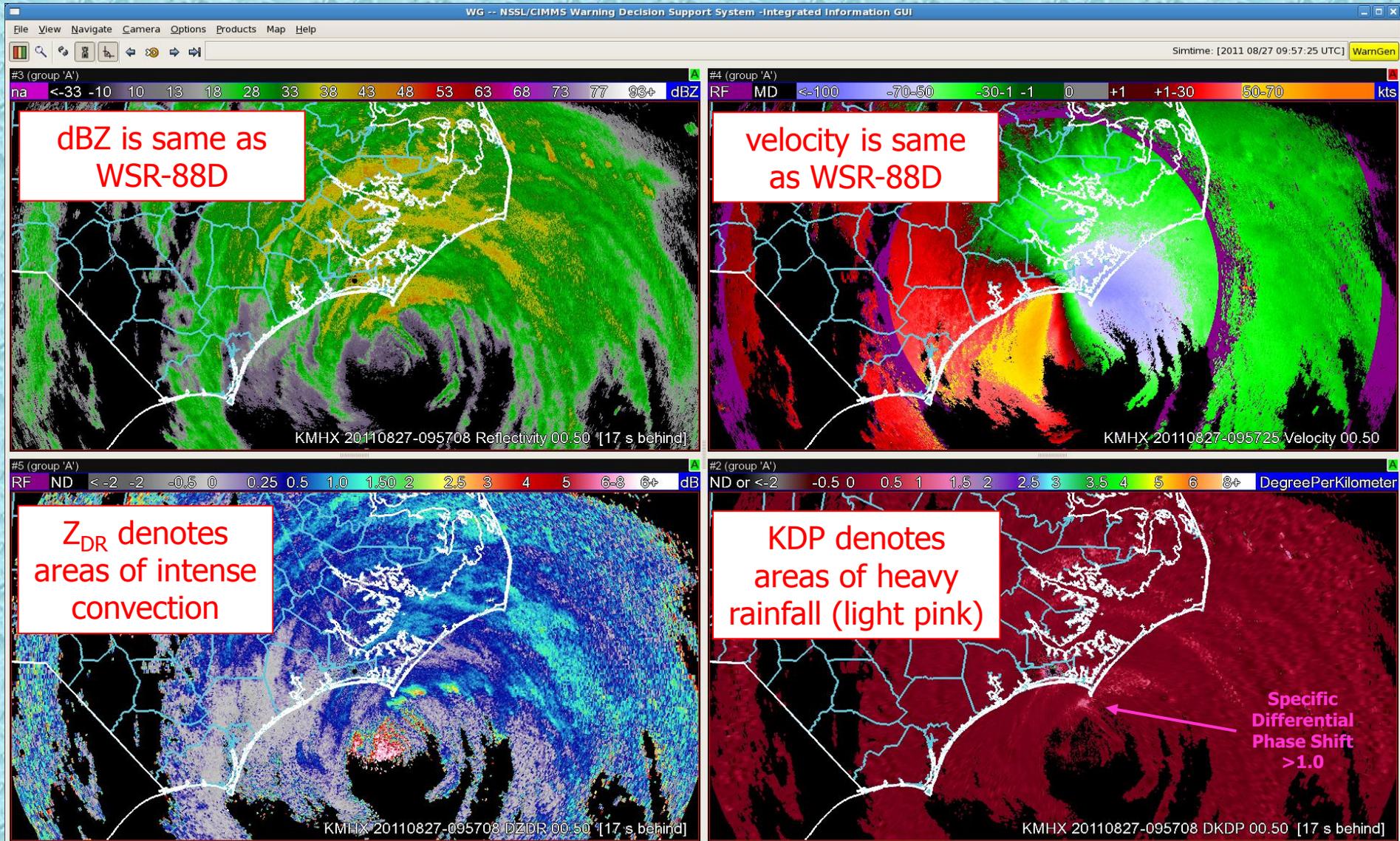


Example: $Z_h = 100,000$ (i.e. 50 dBZ) and $Z_v = 317,000$ (i.e. 55 dBZ)

$$\begin{aligned}\text{Therefore, } Z_{DR} &= 10 \log (Z_h/Z_v) \\ &= 10 \log (100000/317000) \\ &= 10 \log (0.315) \\ &= 10 \times -0.501 \\ Z_{DR} &= -5.01\end{aligned}$$

- Differential reflectivity values above 2 dB are commonly observed in rain.
- Although hailstones are not necessarily spherical, studies have shown that they fall with a tumbling motion -- meaning a field of falling hailstones within the radar resolution volume will "appear" to consist of nearly spherical hydrometeors. Therefore, the value of dBZ_{DR} for hail is usually close to zero.
- Some graupel and hail hydrometeors with a conical shape can fall with their major axes oriented in the vertical. In these cases, the dBZ_{DR} will be found to be negative.

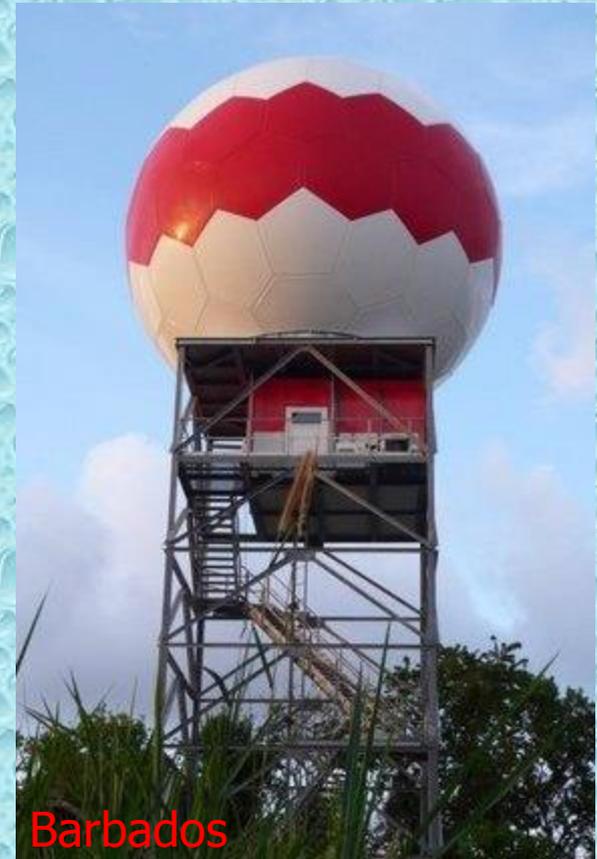
Dual-Pol Radar Example -- Hurricane Irene (2011)



Provides better Z-R relationship and rainfall estimates by determining precipitation type

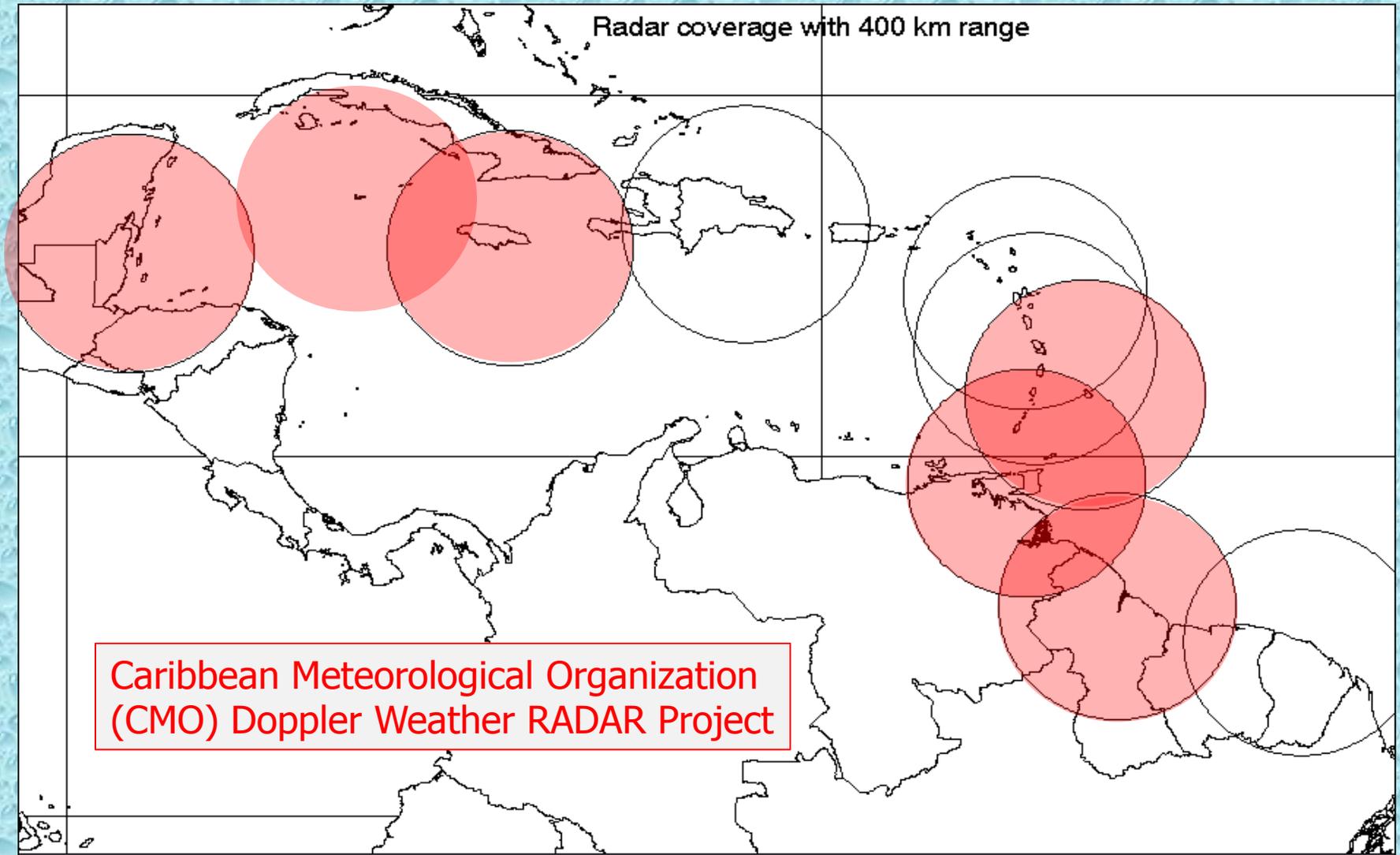
Caribbean Meteorological Organization (CMO) Doppler Weather RADAR Project

- €13-million (Euro) project that aims to replace old radars with new Doppler radars.
- Goal: provide improved awareness of approaching tropical cyclones and heavy rainfall events.
- New radars installed in Barbados, Belize, Guyana, and Trinidad.



<http://www.cmo.org.tt/radar.html>
<http://www.cmo.org.tt/management.html>
http://www.cdera.org/workshop/un-spider/day2/LAYNE-CMO_Presentation.pdf
www.cmo.org.tt/docs/CMC46/CMC46_Docs/CMC46_Doc_8.doc

Radar coverage with 400 km range

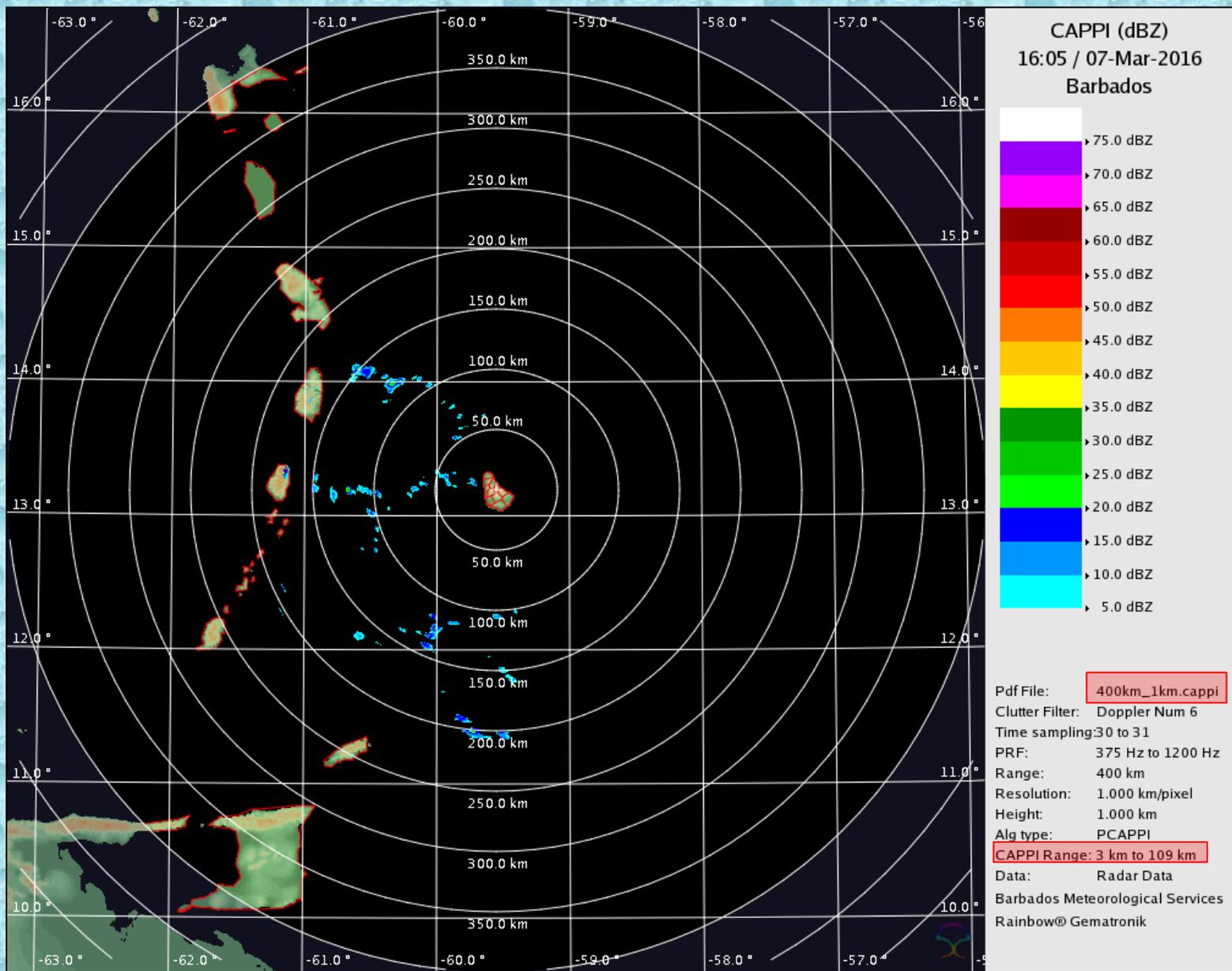


Caribbean Meteorological Organization
(CMO) Doppler Weather RADAR Project

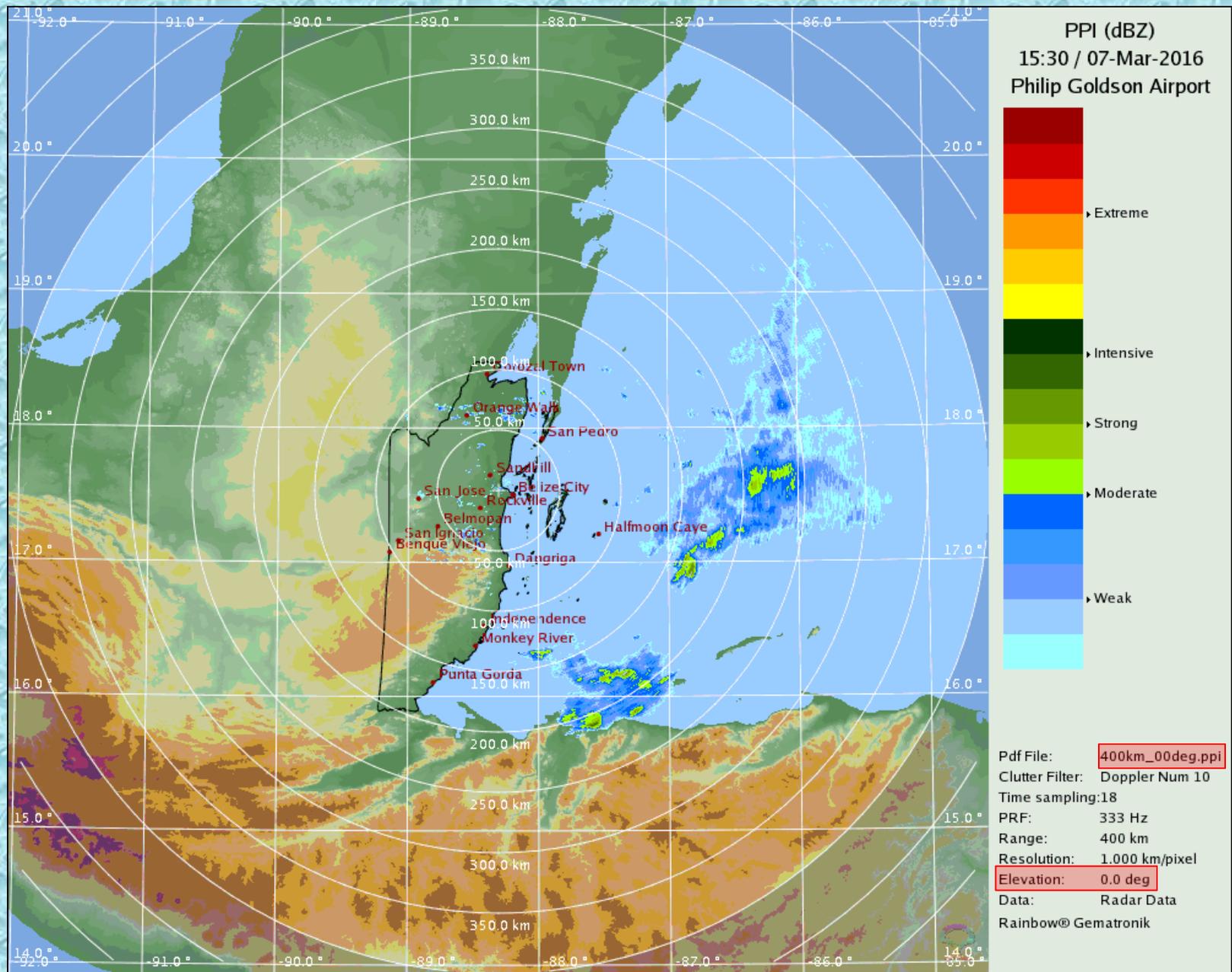
SELEX-Gematronik was awarded the contract for the four METEOR 500S S-Band weather radar systems

- ❑ S-band/10 cm wavelength for maximizing precipitation detection
- ❑ 8.5-meter diameter parabolic antenna dish
- ❑ Not dual-polarization capable
(NOTE: 700S model does have dual-polarization capability)
- ❑ $\leq 1.0^\circ$ beamwidth

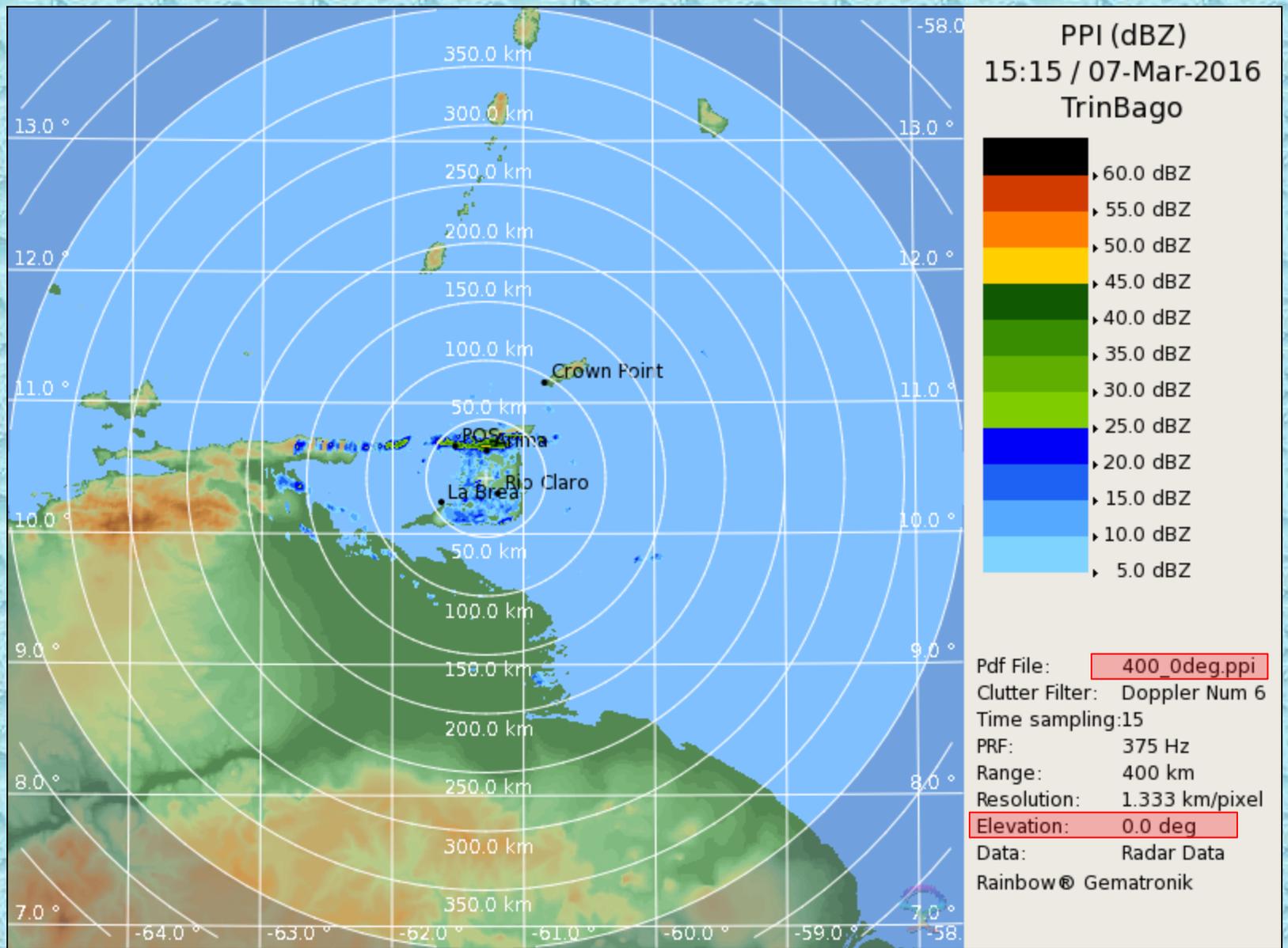




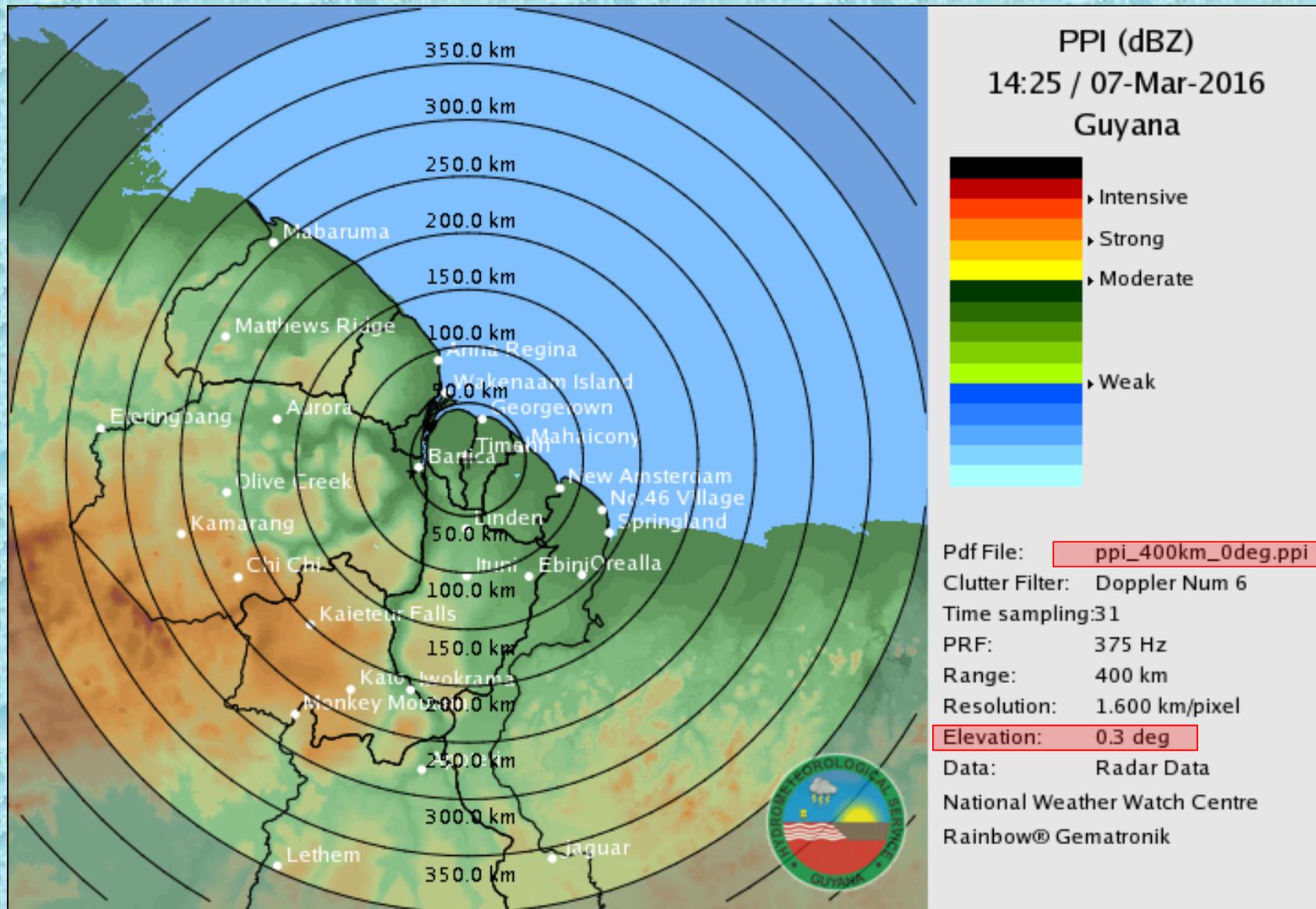
Barbados Doppler Weather Radar



Belize Doppler Weather Radar



Trinidad & Tobago Doppler Weather Radar



Guyana Doppler Weather Radar

Conclusions

- Wavelength is a key design consideration for weather radar (WSR-88D uses 10 cm)
- Doppler velocity is an effective tool in determining tropical cyclone intensity, and detecting rapid intensification.
- Future – dual polarization will give better precipitation estimates.
- NEXRAD-in-space will generate radar data for the entire Atlantic basin.

